

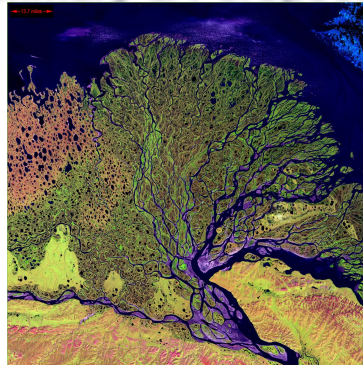
# Representing natural land surface heterogeneity in Numerical Weather Prediction and Earth System models: Benefits and Limitations of current schemes in presence of large contrasts

*presented by*

**Gianpaolo Balsamo**

**Outline:**

Introduction: about models and information cycle  
TESSEL LSM scheme and its evolution  
forest/snow contrasts  
forest/lakes contrasts  
complexity: when does it pay off?  
Summary & Perspectives



**Acknowledgements:**

**Anton Beljaars, Andrea Manrique-Sunen, Souhail Boussetta, Emanuel Dutra,  
Clement Albergel, Patricia de Rosnay, Anna Agusti-Panareda, Thomas Haiden**

## Abstract

Natural and anthropogenic variability that characterizes the Earth's surface is very much a driver for **complexity** and increased resolution in NWP and ESS models, with existing or upcoming **support** from **Earth-Observation datasets** such as those derived from MODIS, SPOT, PROBA-V, Sentinels ESA's program satellites, all attaining sub-kilometer remote-sensing capabilities. While a higher **resolution** improves the description of contrasting surfaces such as land-water or snow-forest, large part of these **contrasts** is bound to remain a **sub-grid** parameterization's issue due to fractal nature of the surface: coastlines, rivers, lakes, forests, urban-areas are rarely pure land-use categories even when approaching kilometric scales. Representing those contrasts is important for partitioning energy and water fluxes but also for carbon-exchange and therefore have received attention both in NWP and ESM. The **land surface tiling** is often a method used for such purposes as it can accommodate several parameterizations for the different surface types. We have examined the capacity of the tiling to represent large natural contrasts in presence of snow, forests and lakes, in the framework of the land surface scheme operationally used at ECMWF. The **benefits** on representing the fluxes and the **limitations** coming from lateral decoupling (lack of mixing) are highlighted in a set of field-site examples and the impact is evaluated in global sensitivity experiments. Preliminary ideas to go beyond the tiling concepts will be discussed.

# Models: projection of our understanding

## How do we learn?

The information cycle (from *E. Morin, "On complexity"*):

● **Information → Knowledge → Wisdom**

In the scientific world:

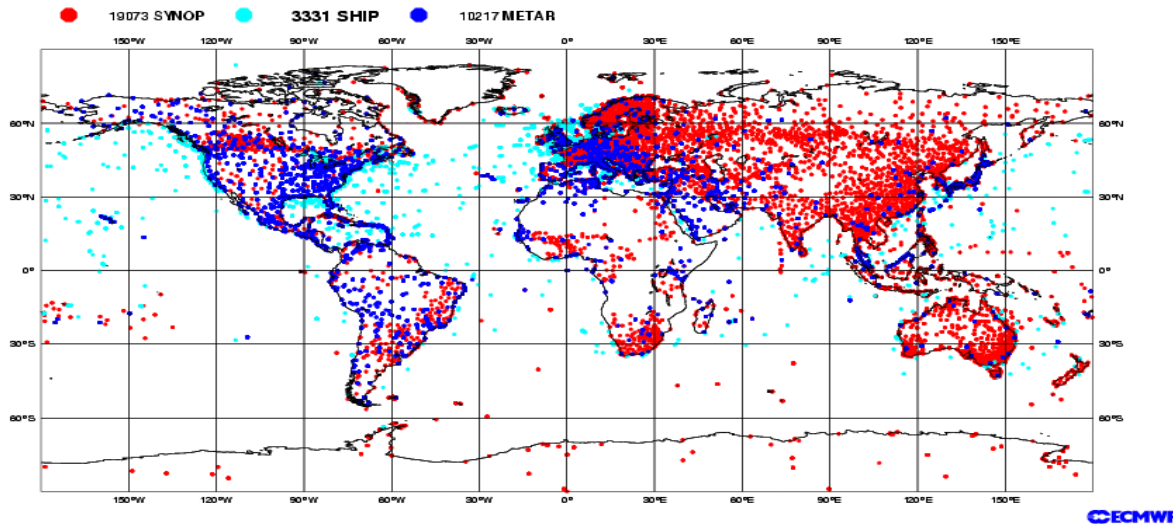
● **Data & Obs → Model → Monitor & Prediction**

In the decision-makers world:

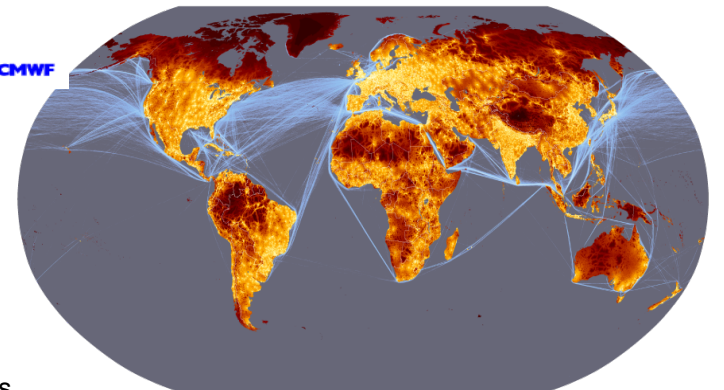
● **Products → Cost-Loss → Action & Planning**

# The conventional surface observations

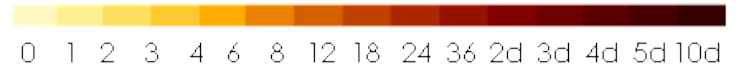
ECMWF Data Coverage (All obs DA) - Synop-Ship-Metar  
10/May/2012; 00 UTC  
Total number of obs = 32621



Surface based conventional observation networks are sparse and heterogeneous in coverage.



Travel time to major cities  
(>50k inhabitants)  
expressed in hours/days.  
Source:JRC



# Land surface satellite remote sensing

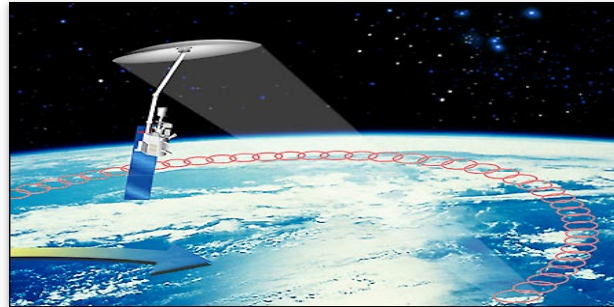
Great potential to have a more homogeneous coverage of the land surface...several challenges in extracting information content

## L-band Tb

SMOS ESA mission

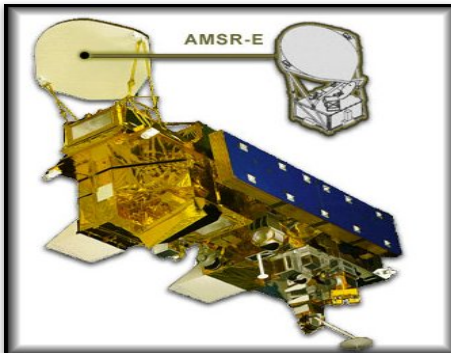


SMAP (Soil Moisture Active Passive) NASA mission



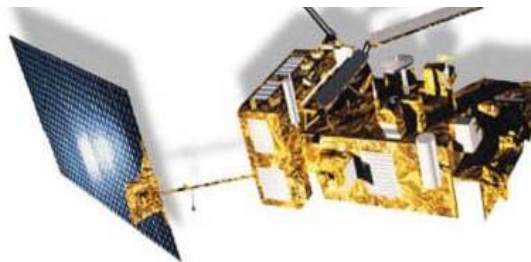
## C-band passive Tb

AQUA AMSR-E instrument  
(05/2002)



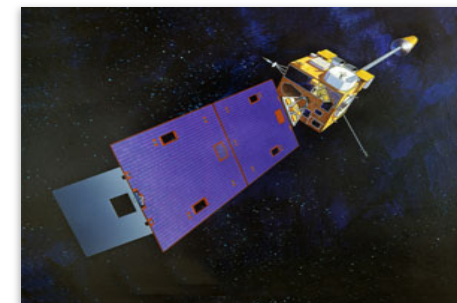
## C-band active

MetOP ASCAT (2008- )

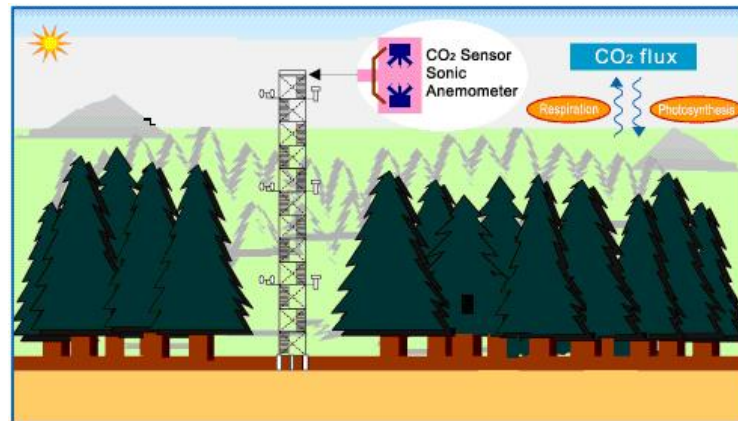
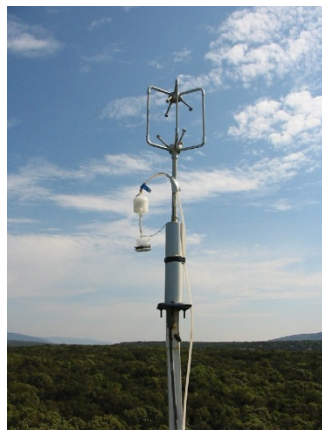
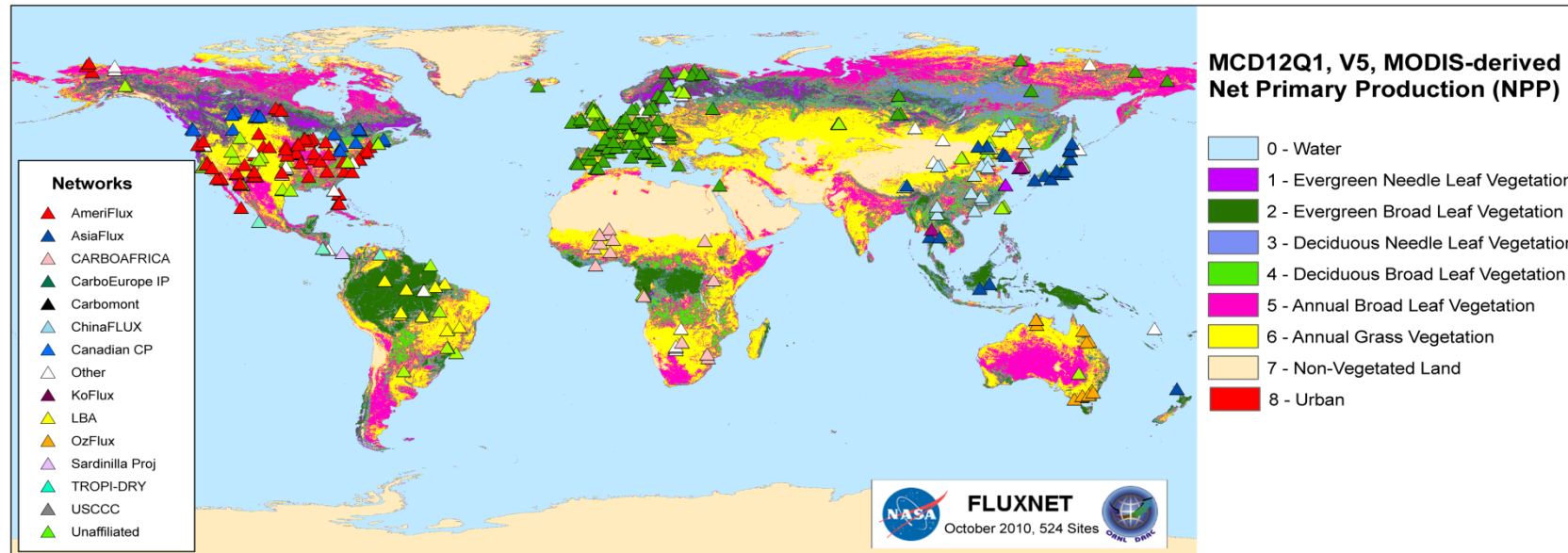


## IR Ts

GOES-E GOES-W  
MSG MTSAT



# FLUXNET: An example of research network

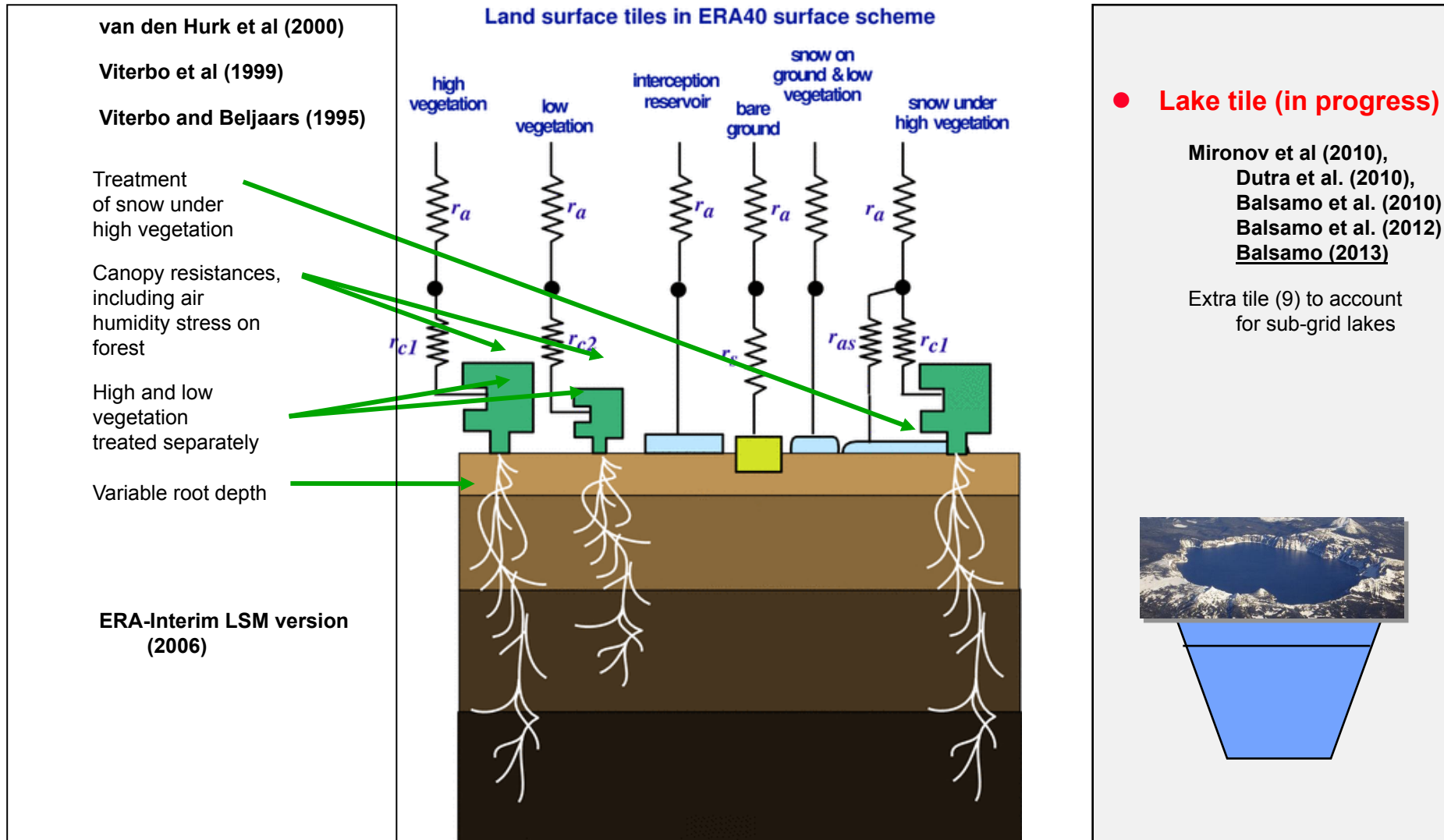


The in-situ research networks are also fundamental for “fit-for-purpose” model development.

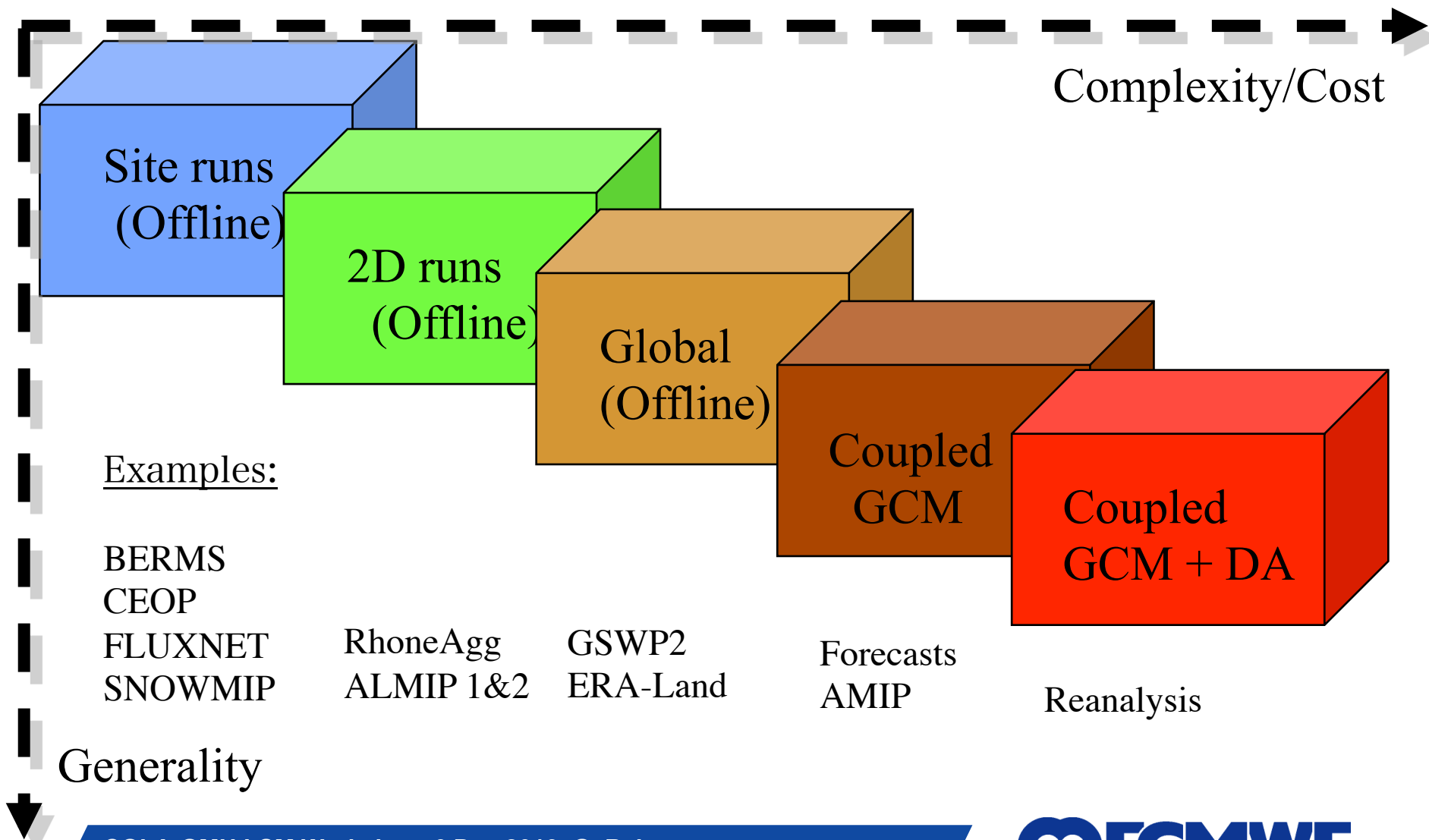
Another good example is the ISMN for soil moisture (maintained by TU-Wien)

# The ECMWF TESSEL scheme (main tiles)

- Tiled ECMWF Scheme for Surface Exchanges over Land



# A step-wise approach to model complexity





# Integration of the offline surface system

Driving the surface offline with imposed meteorological forcing is a practical way to break the complexity of the Earth system and it has enabled R&D at reduced computational cost. This is fully integrated in ECMWF software (PrepIFS/Xcdp/MARS) and has enabled the generation of **ERA-Interim/Land**

The image illustrates the integration of an offline surface system. On the left, a hand holds a paper sun over a field, with red arrows pointing from the sun to the ground, representing imposed meteorological forcing. A dashed black line separates the sky from the ground. On the right, there are three overlapping screenshots of software interfaces. The top one is a 'Create new experiment ID' dialog box with a grid of checkboxes; the 'surface' checkbox is circled in red. The middle one is a file tree view for an experiment named 'foyd', showing a 'surface' folder with sub-items like 'make', 'initsurf', 'forcing', 'main', and 'lag'. The bottom one is an 'INLAND specification' dialog box with various options and dropdown menus.

Balsamo et al. (2013), ERA-Interim/Land: <http://www.hydrol-earth-syst-sci-discuss.net/10/14705/2013/hessd-10-14705-2013.html>

# Land surface model evolution since ERA-I

2007/11	2009/03	2009/09	2010/11	2011/11	2012/06
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● **Hydrology-TESEL**

Balsamo et al. (2009)  
van den Hurk and Viterbo (2003)

Global Soil Texture (FAO)

New hydraulic properties

Variable Infiltration capacity & surface runoff revision

● **NEW SNOW**

Dutra et al. (2010)

Revised snow density

Liquid water reservoir

Revision of Albedo and sub-grid snow cover

● **NEW LAI**

Boussetta et al. (2011)

New satellite-based Leaf-Area-Index

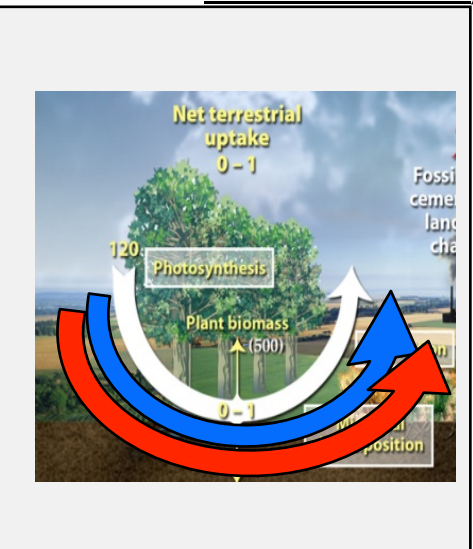
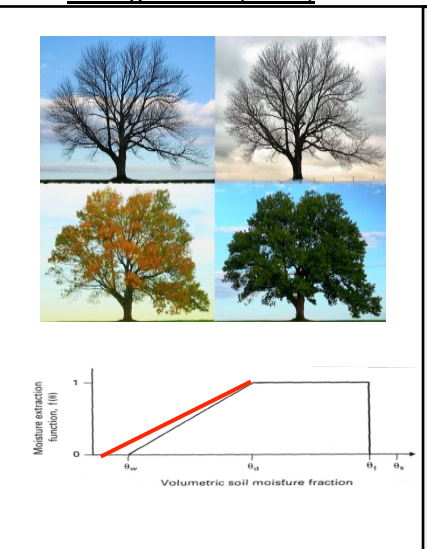
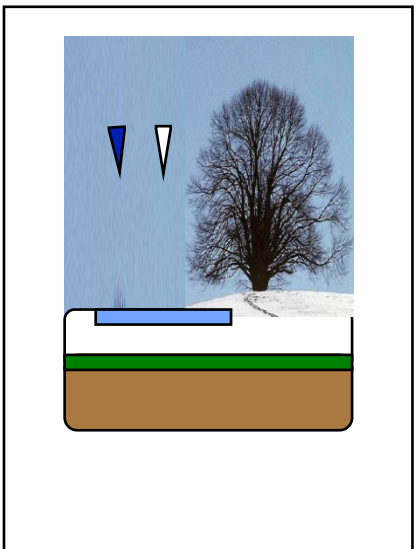
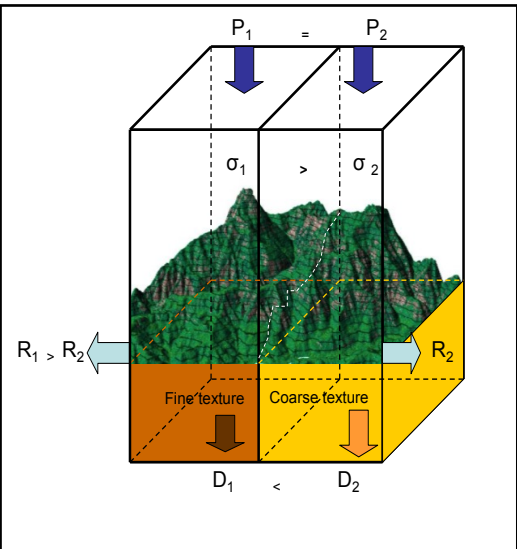
● **SOIL Evaporation**

Balsamo et al. (2011),  
Albergel et al. (2012)

● **H<sub>2</sub>O / E / CO<sub>2</sub>**

Integration of Carbon / Energy / Water cycles at the surface (GEOLAND-2 based & GMES funded)

Calvet et al. (1998)  
Jarlan et al (2007)  
Boussetta et al. (2013a,  
Boussetta et al. 2013b)



# Representing land surface contrasts

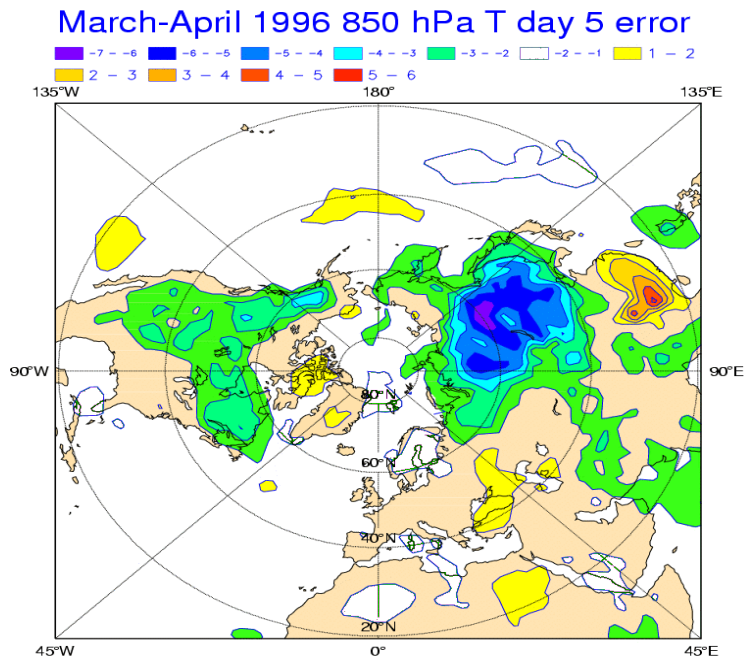
- **Snow-Forest**

- **Lakes-Forest**

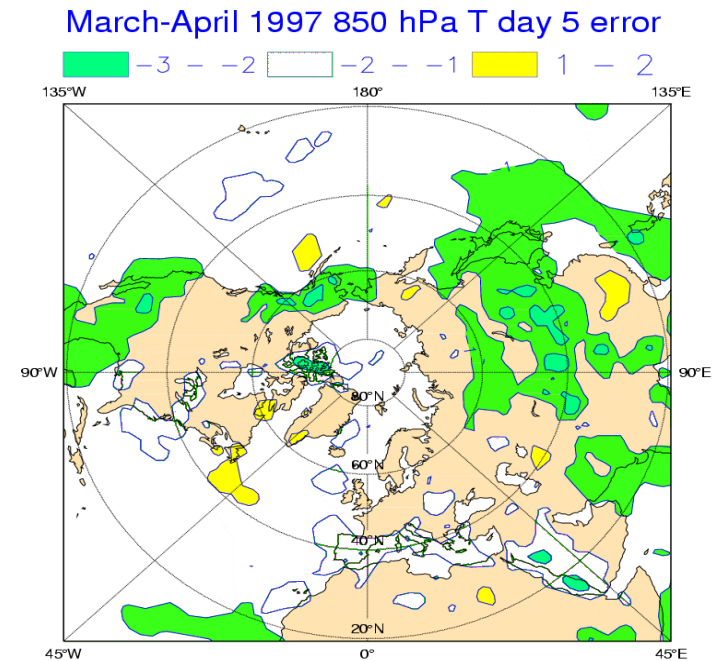
# Why Forest & Snow contrasts matters?

A lower albedo of snow+forest tile in the boreal forests (1997) reduced dramatically the spring (March-April) error in day 5 temperature at 850 hPa

1996 operational bias



1997 operational bias



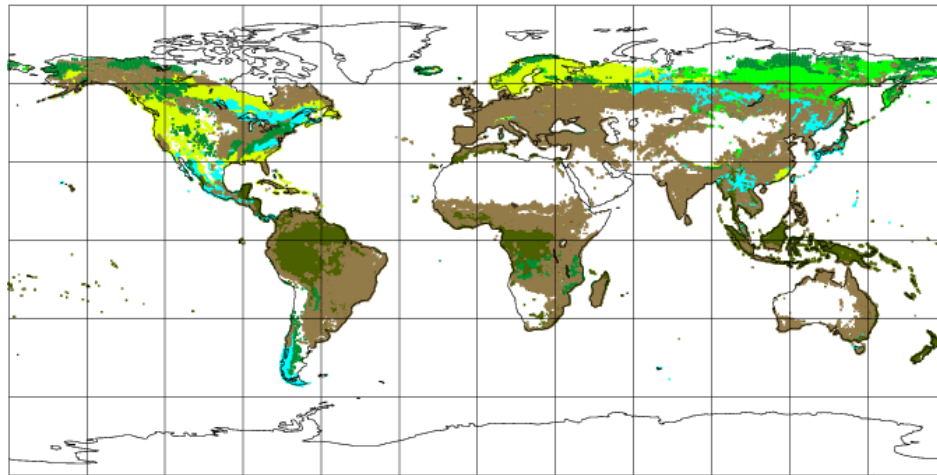
Viterbo and Betts, 1999

# Representation of Vegetation (forest)

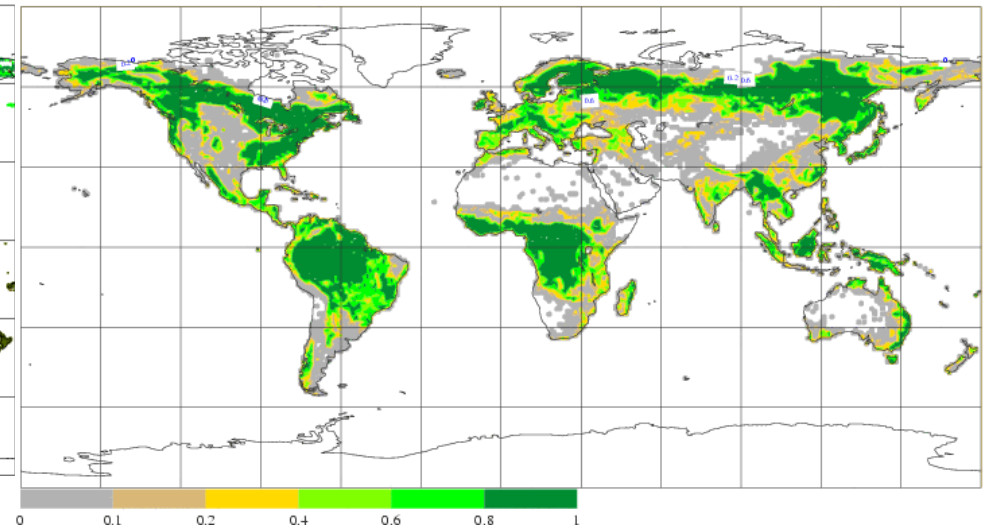
The dominant forest type and its cover are prescribed by GLCC v1.2 global static map (Loveland et al. 1998)

## FOREST TYPE

ever needle   deci needle   deci broad   ever broad   mix forest   int forest



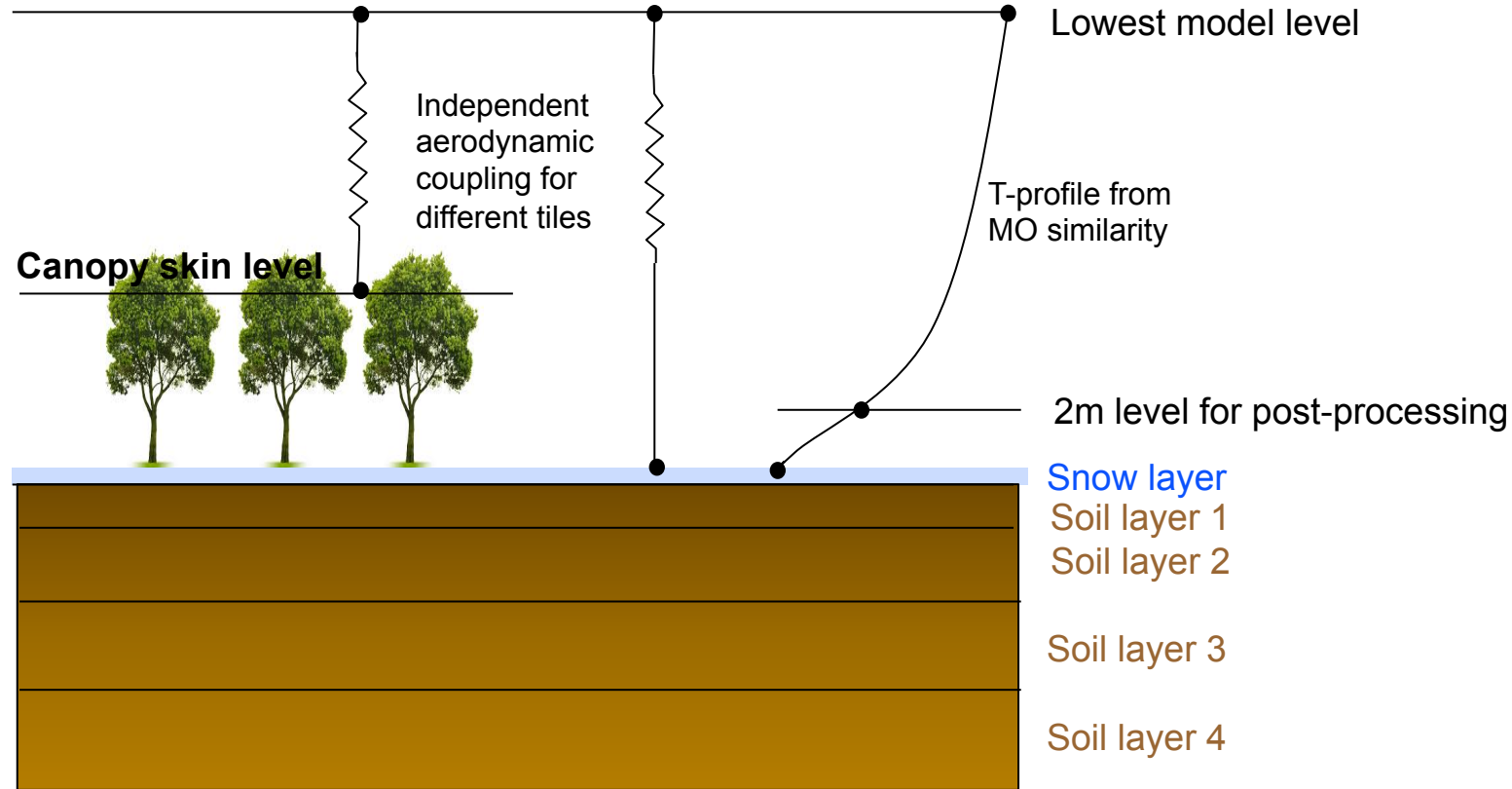
## FOREST COVER FRACTION



Aggregated from GLCC 1km

# Can we simulate snow-forest T contrasts?

(i) tile with snow under vegetation and (ii) tile with exposed snow

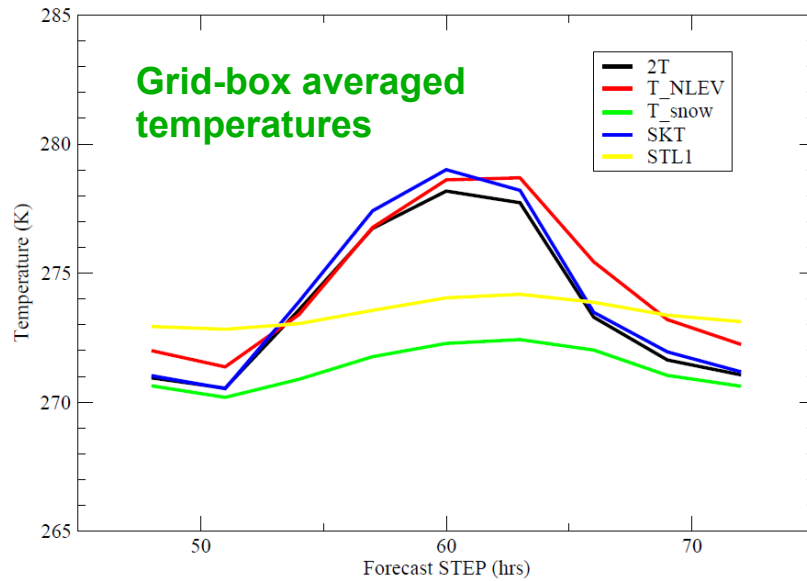


Even if the forest is dominant, the vertical interpolation to the 2m level is done for the exposed snow tile (SYNOP stations are always in a clearing).

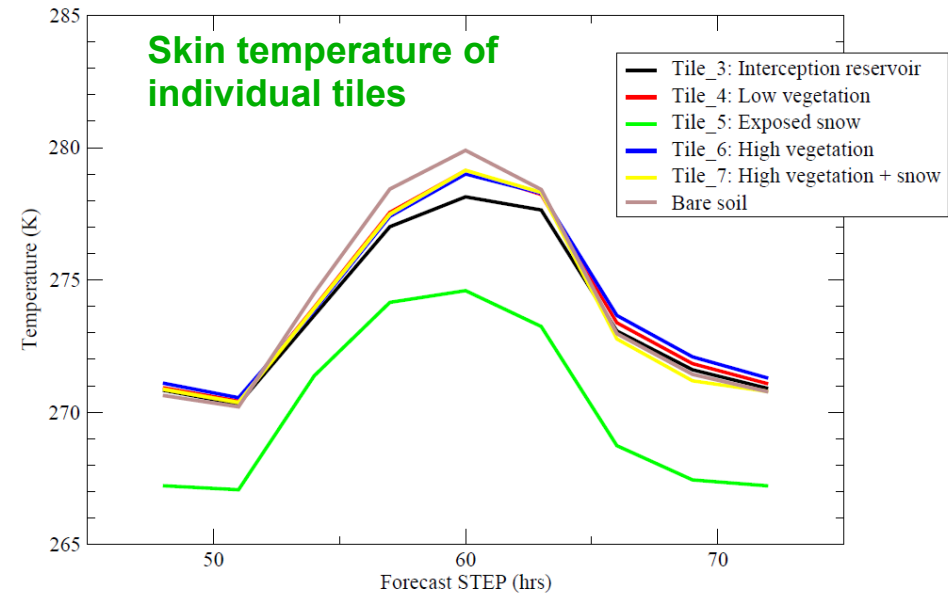
During day time, the forest heats the atmosphere. At sunset exposed snow tile becomes very stable cutting off turbulent exchange. Therefore snow temperature and T2 drop too much. In reality forest generated turbulence will maintain turbulent exchange over the clearing and prevent extreme cooling.

# Tiles temperature split is too strong

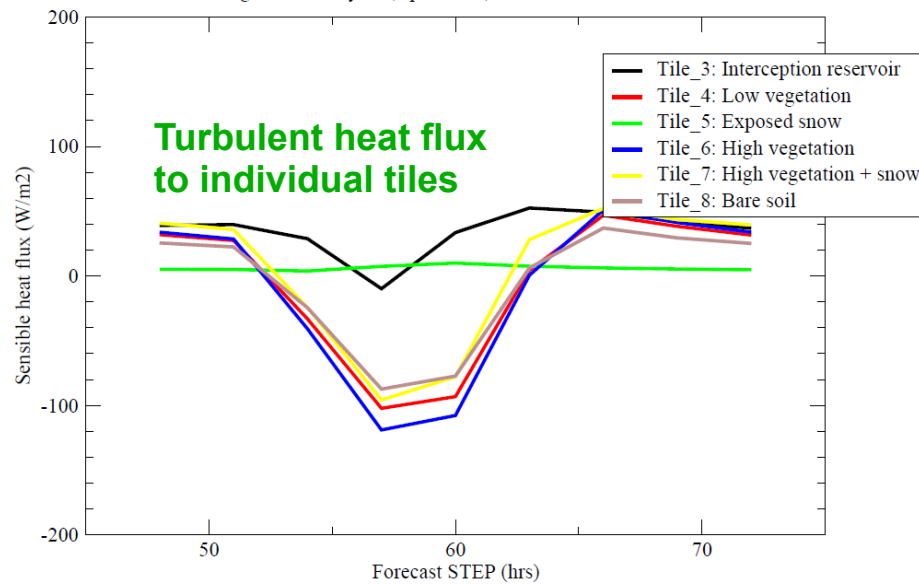
Averaged diurnal cycle (April 2013): Tile averaged temperatures



Averaged diurnal cycle (April 2013): Tiled skin temperatures

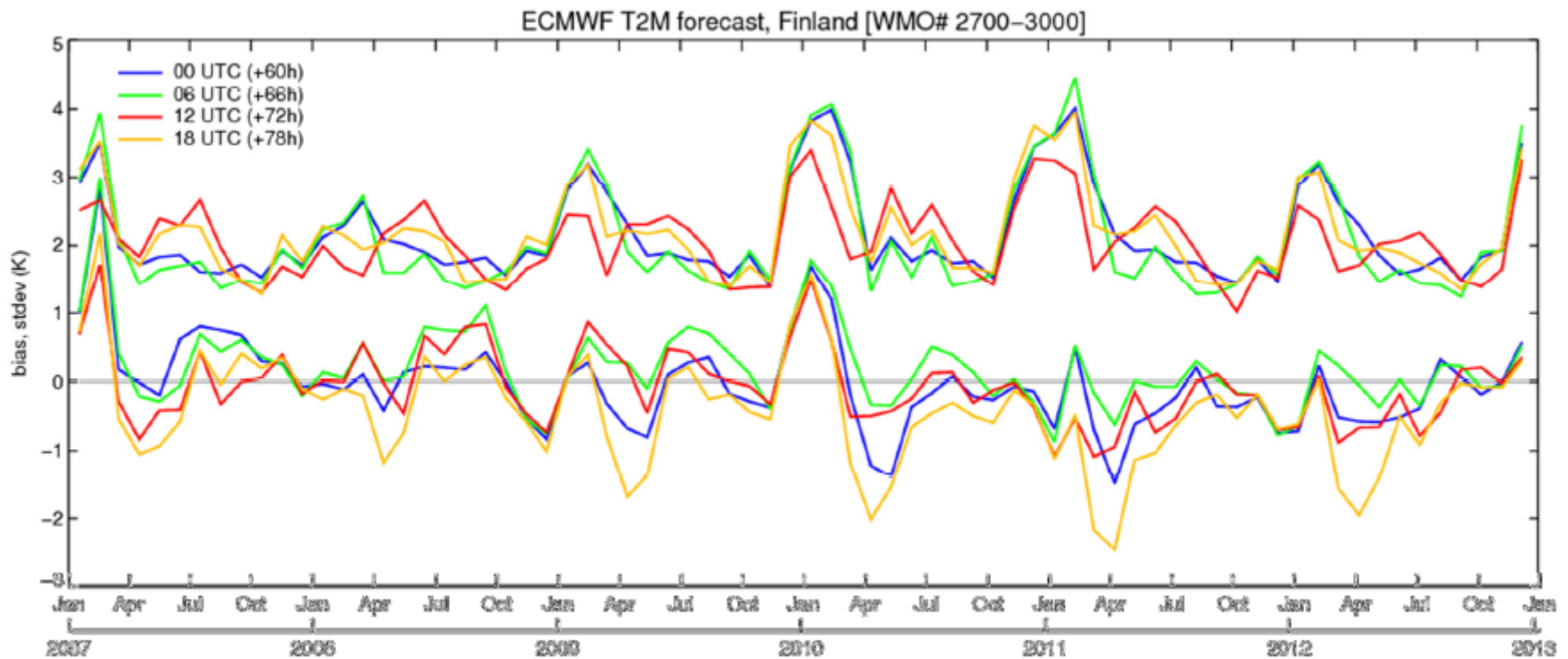


Averaged diurnal cycle (April 2013): Tiled sensible heat fluxes



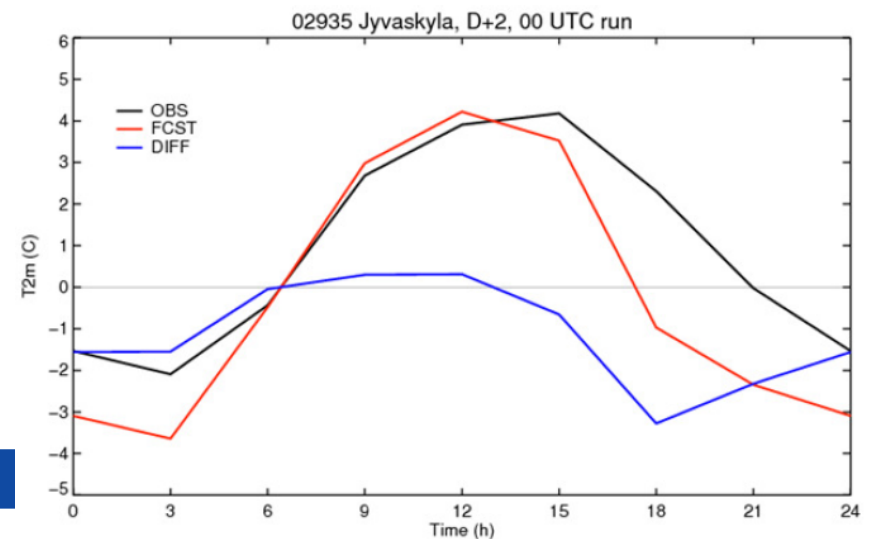
The heat flux towards the exposed snow is nearly zero! This is unrealistic and show a shortcoming of tiling. Do we see it in NWP scores?

# Spring temperature biases over Scandinavia



Scandinavian countries show a spring time cold bias mainly at 18 UTC related to snow melt in forested areas. The bias has a distinct diurnal cycle.

by Thomas Haiden



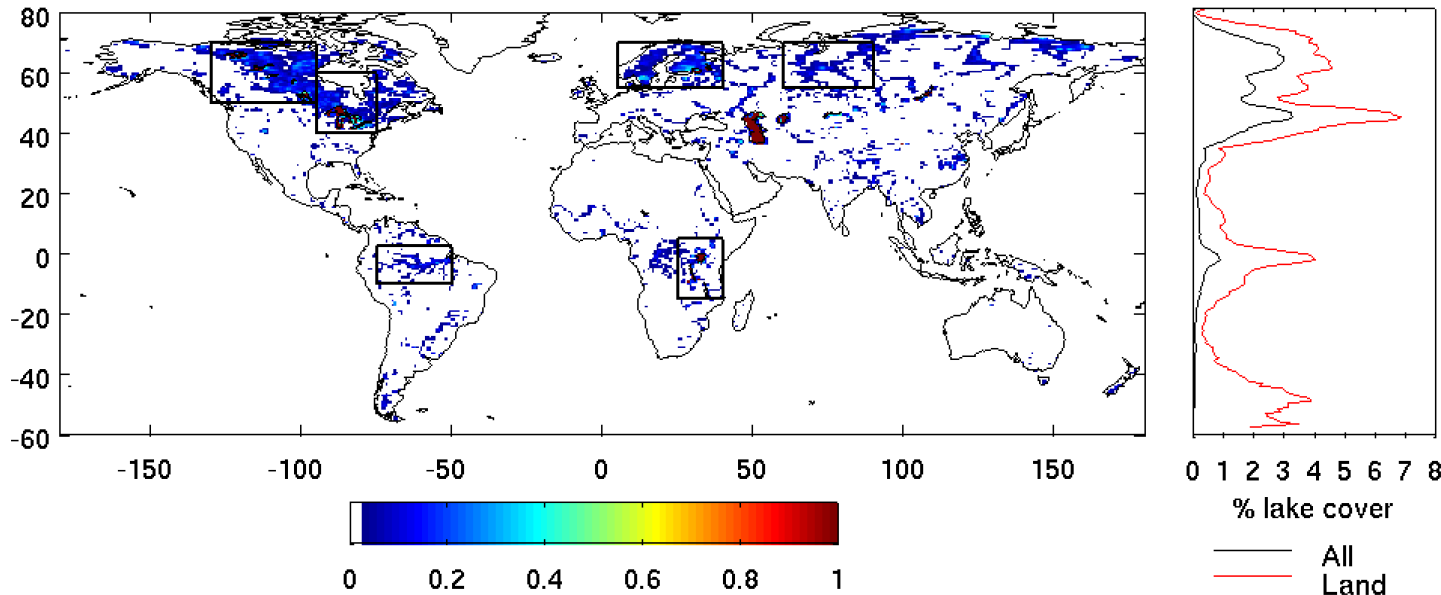


# Representing lakes in LSM: Motivations

E. Dutra, V. Stepaneko, P. Viterbo, P. Miranda, G. Balsamo, 2010 BER

Motivation: a sizeable fraction of land surface has sub-grid lakes

## LAKE COVER FRACTION

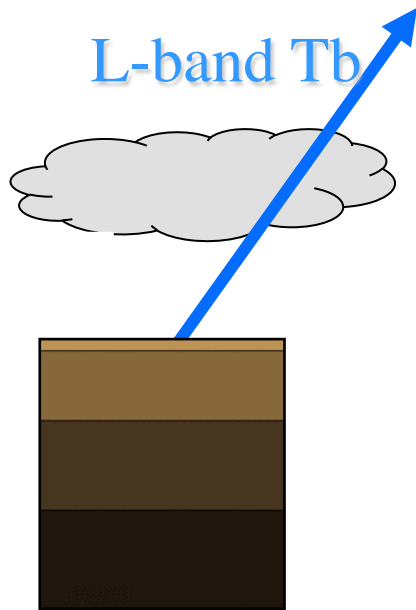


N° Points  $0.05 < C_{lake} < 0.5$

<b>Canada</b>	309/754 <b>41%</b>
<b>USA</b>	175/482 <b>36%</b>
<b>Europe</b>	170/385 <b>44%</b>
<b>Siberia</b>	104/467 <b>22%</b>
<b>Amazon</b>	81/629 <b>13%</b>
<b>Africa</b>	74/584 <b>13%</b>

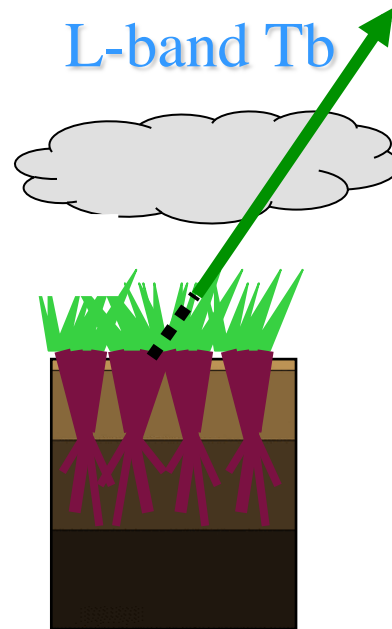
Microwave channels (and L-band in particular) is very sensitive to lake presence (even 1-2%) and lake temperature: SMOS/SMAP forward modelling impact

# Microwave Remotely sensing from space: Relevance of open-water in forward modelling



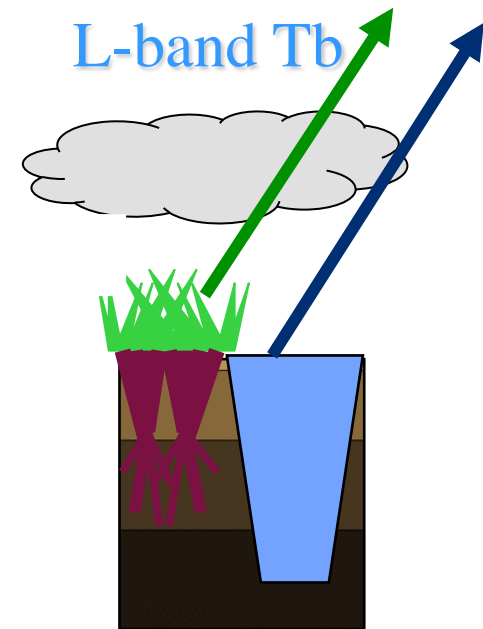
Soil moisture modifies soil dielectric constant  $\rightarrow$  emissivity  $\epsilon$

$$T_{b\_soil} = \epsilon T_s$$



Vegetation attenuates soil emission + emits its own TB

$$T_b \text{ influenced by vegetation layer [f(LAI)]}$$



Lakes create a strong cold signal, masking the signal of soil

$$T_b \text{ varying with lake temperature [f}(T_{skin})]$$

L-band Tb  
C-band Tb

Sounding soil depth	Frequency	Wavelength	Atmospheric absorption
~5 cm	1.4 GHz	21 cm	<b>Negligible</b>
~1cm	6.9 GHz	5 cm	Low (except rainy area)

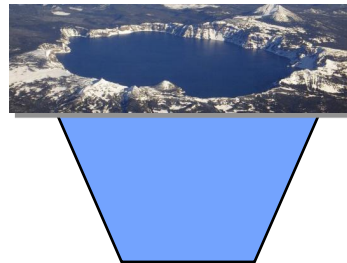
# Can we simulate Forest and Lakes contrasts?

Andrea Manrique-Sunen et al (2013, JHM)

**Meteorological forcing:** ERA-Interim reanalysis

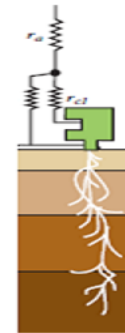
Model was run for the year 2006, doing 3 iterations

LAKEHTESSEL



Lake: Full coverage of inland water  
Lake depth = 4 m  
Water extinction coefficient = 3 m<sup>-1</sup>

HTESSSEL



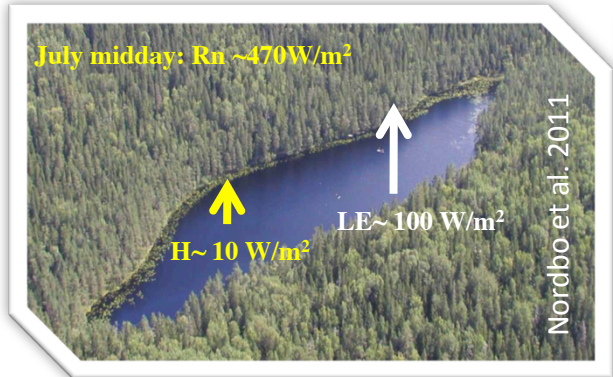
Forest: Full coverage of high vegetation  
Vegetation type: Evergreen needleleaf trees  
Soil type: Medium texture

$$\text{Energy balance in the surface } R_n + SH + LE = G$$

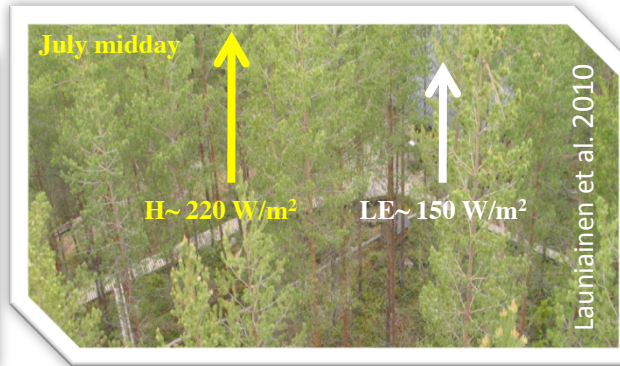
# A Finnish lake case study

Manrique-Suñén et al. (2013, JHM)

## LAKEHTESSEL



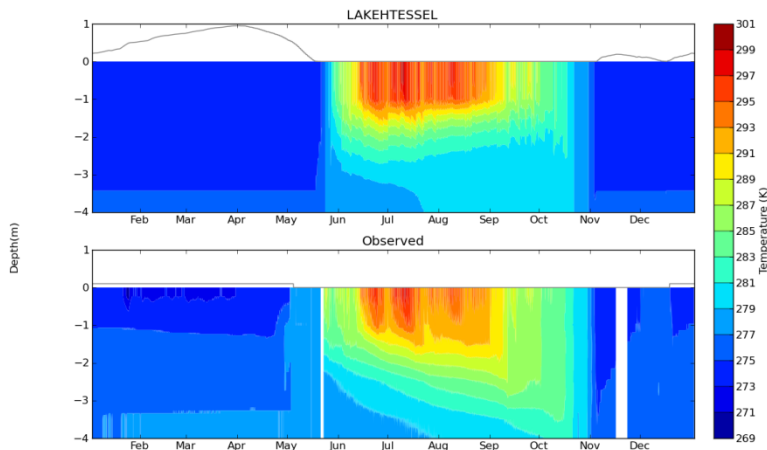
## HTESSSEL



### ● FLake

Mironov et al (2010),  
 Dutra et al. (2010),  
 Balsamo et al. (2010)  
Balsamo et al. (2012)

Extra tile (9) to account  
 for sub-grid lakes

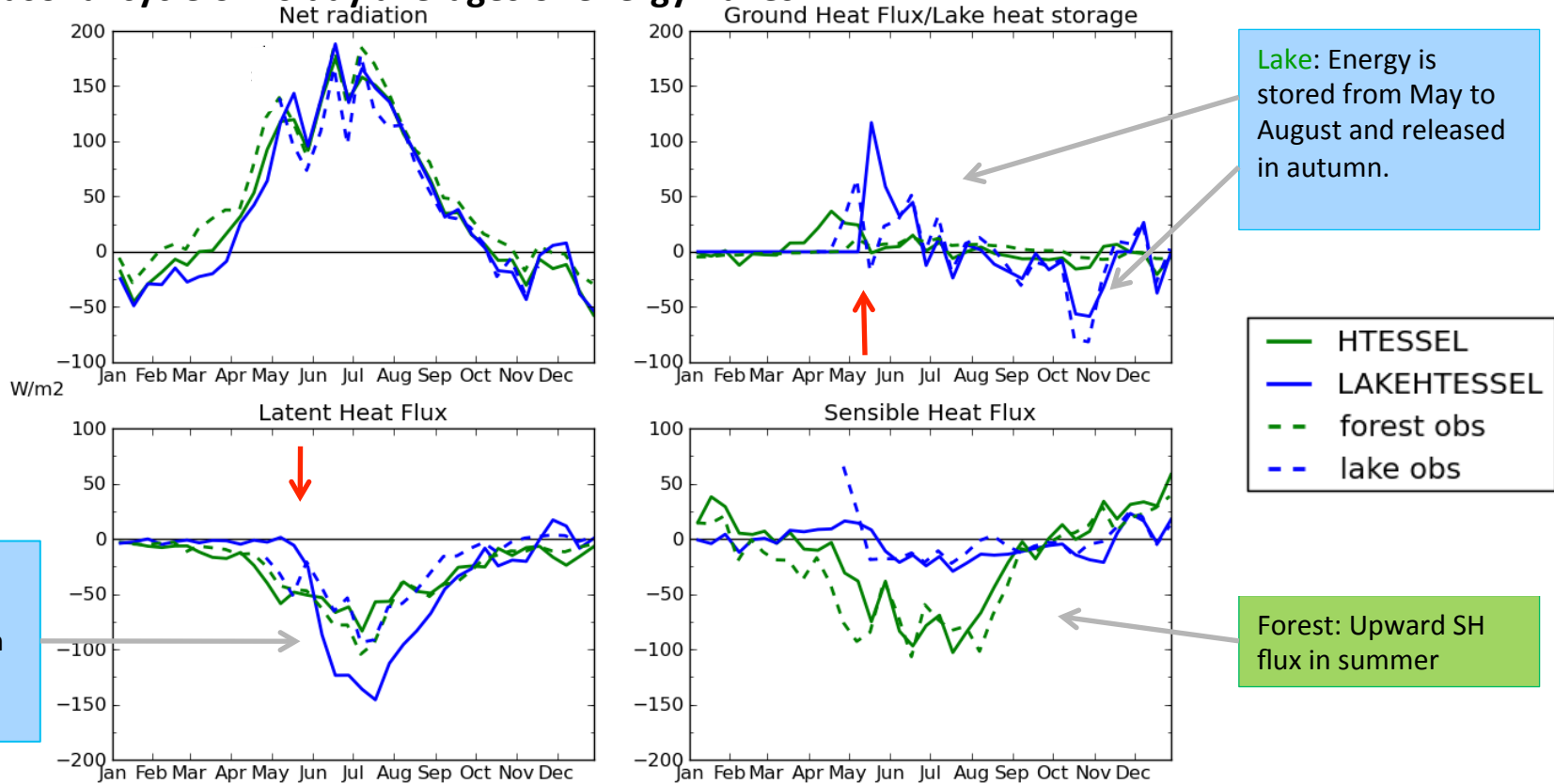


Over a lake specialized site observations can be compared with FLake (Mironov et al. 2010) model output as provided by the LAKEHTESSEL model version.

Collaboration with Annika Nordbo & Ivan Mammarella (U. Helsinki)

# Energy fluxes: Seasonal cycles

Seasonal cycle of 10 day averages of energy fluxes



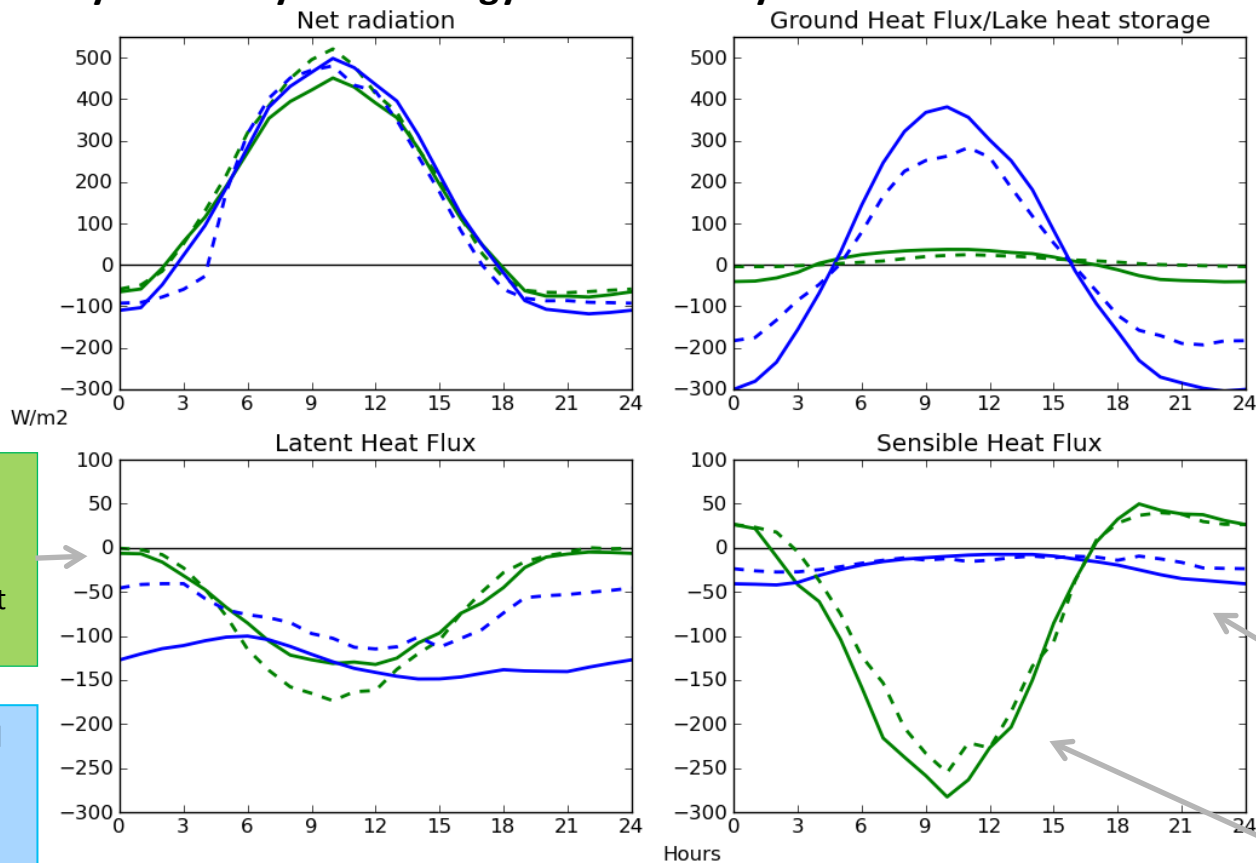
Sign convention: Positive downwards

The timing of the lake's energy cycles is influenced by the ice cover break up, and it is delayed by 14 days in the model

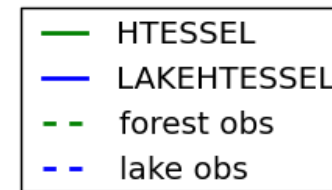
Main difference between both sites is found in the energy partitioning into SH and G

# Energy fluxes: Diurnal cycles

Monthly diurnal cycle of energy fluxes for July



Very good representation by the model of diurnal cycles and particularities of each surface



Forest evaporation is driven by vegetation, so it is zero at night

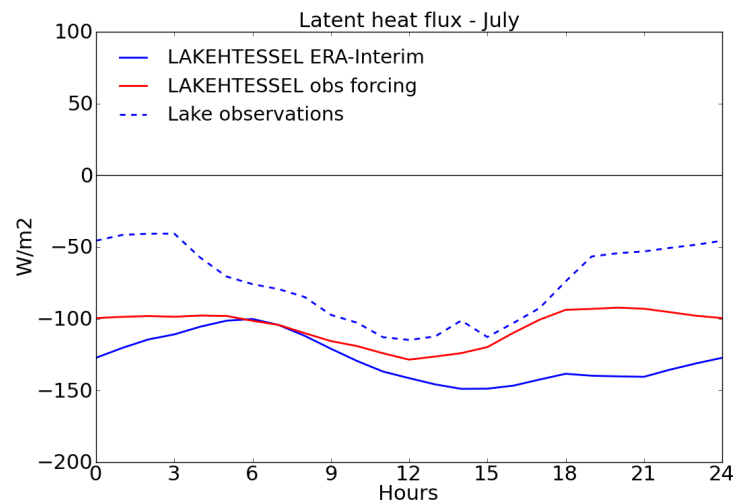
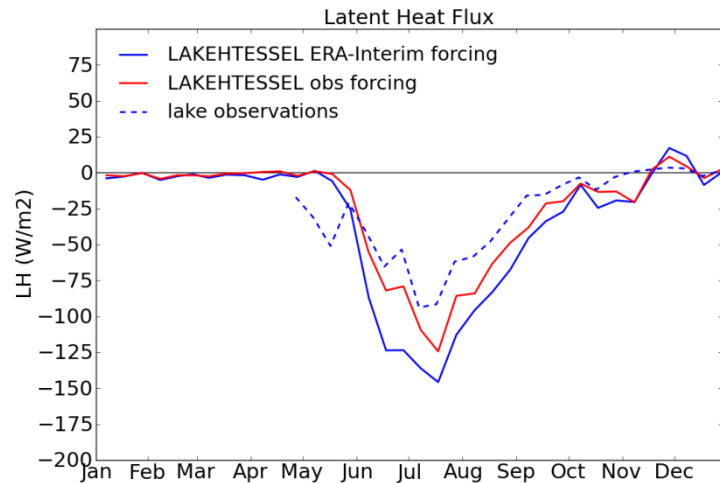
Lake LH diurnal cycle: overestimation in evaporation

Lake SH maximum is at night

Forest SH maximum is at midday

Main difference between both sites is found in the energy partitioning into SH and G

# Use of observed forcing vs ERA-Interim for the lake site: highlight tiling shortcoming



Seasonal cycle:

The use of observed forcing reduces the RMSE in evaporation from  $32 \text{ W m}^{-2}$  to  $19 \text{ W m}^{-2}$

Diurnal cycle for July:

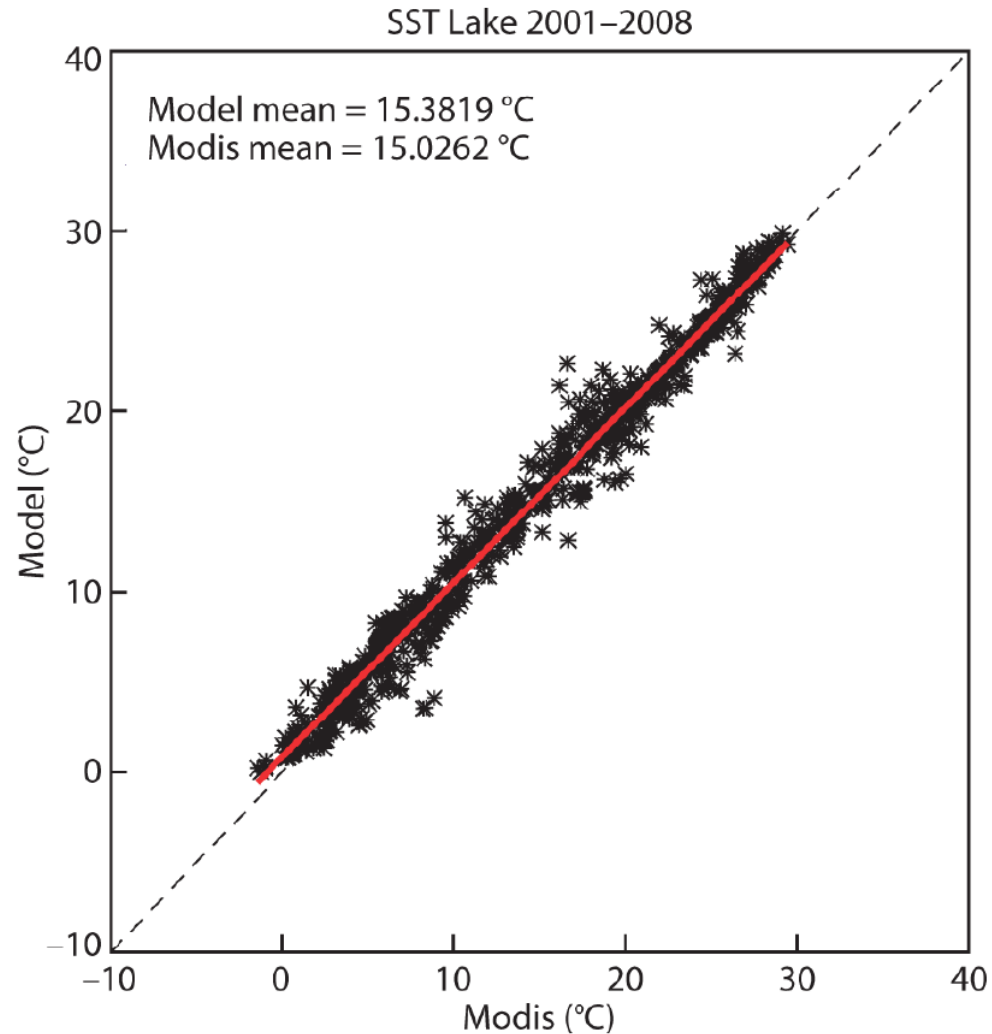
The evaporation is reduced, but errors remain at night.

The model's transfer coefficients might not be appropriate for a calm situation

Manrique-Suñén et al. (2013, JHM)

# Lakes surface temperature (global validation)

Balsamo et al. (2012, TELLUS-A) and TM 648

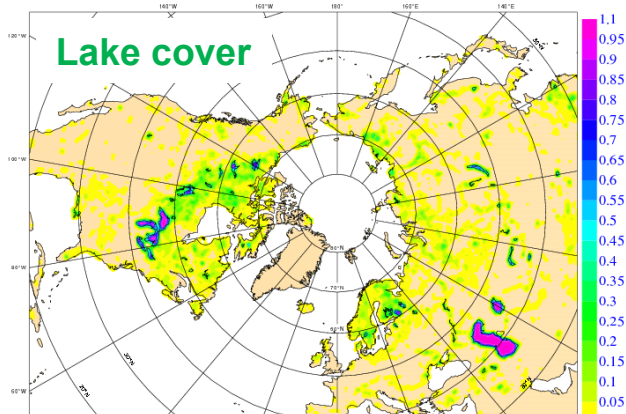


- FLAKE Lake surface temperature is verified against the MODIS LST product (from GSFC/NASA)
- Good correlation  
 $R=0.98$
- Reduced bias  
 $\text{BIAS (Mod-Obs)} < 0.3 \text{ K}$



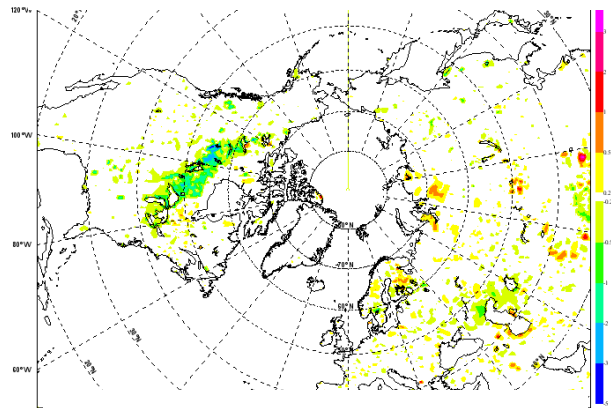
# Impact of lakes in NWP forecasts

Balsamo et al. (2012, TELLUS-A) and ECMWF TM 648



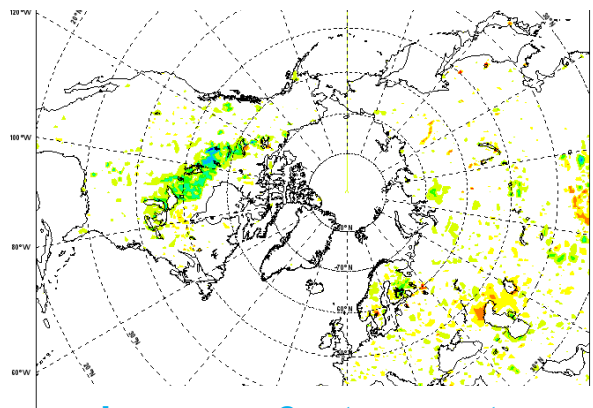
- Forecasts sensitivity and impact is shown to produce a spring-cooling on lake areas with benefit on the temperatures forecasts (day-2 (48-hour forecast) at 2m).

Forecast sensitivity



Cooling 2m temperature  
Warming 2m temperature

Forecast impact



Improves 2m temperature  
Degrades 2m temperature

ERA-Interim forced runs of the FLAKE model are used to generate a lake model climatology which serves as IC in forecasts experiments (Here it is shown spring sensitivity and error impact on temperature when activating the lake model).

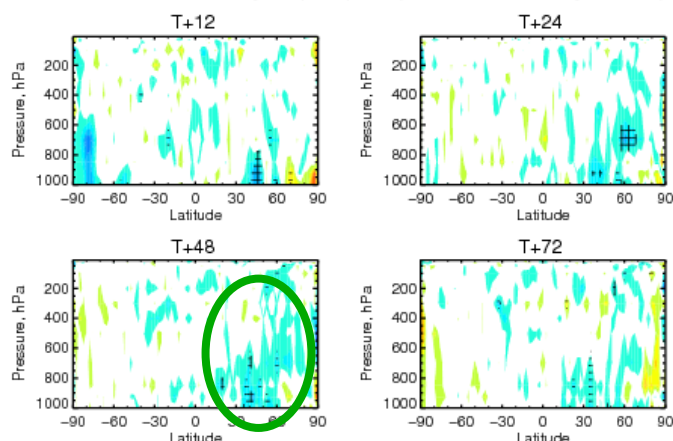
# Implementation of lakes in FC/AN/ENS

## ● AN cycling and initialisation

### Summer experiment

(Temperature scores)

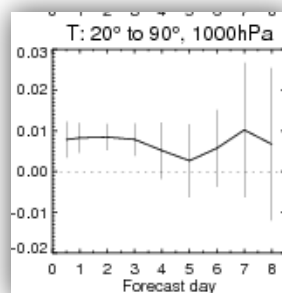
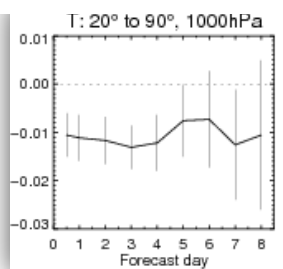
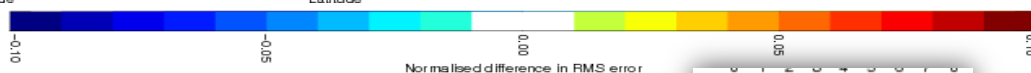
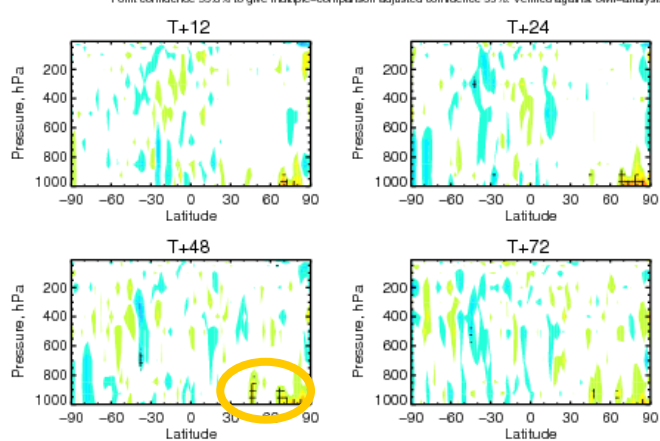
RMS forecast errors in T (fy4o–fy4b), 2–Jun–2012 to 31–Jul–2012, from 52 to  
Point confidence 99.8% to give multiple-comparison adjusted confidence 95%. Verified against own-analysis



### Winter experiment

(Temperature scores)

RMS forecast errors in T (fy4p–fy09), 1–Jan–2013 to 21–Feb–2013, from 44 to 5  
Point confidence 99.8% to give multiple-comparison adjusted confidence 95%. Verified against own-analysis.



- Modelling transitions of lake open water to lake-ice is very challenging and may require a careful initialisation
- Sea-ice is probably in a similar situation (predictive skill severely affected by lack of atmospheric predictability in winter)

Balsamo (2013, ECMWF Autumn Newsletter)

*The initial conditions for lakes will play a very important role, and are particularly relevant for lake icing/open water.*

# Complexity calls?

- **Snow timescales and its links to hydrology**
- **Vegetation treatment and its links to carbon cycle**

# Snow: A good call for complexity?

(Dutra et al. 2010 JHM, Balsamo et al. 2011 EC-NL)

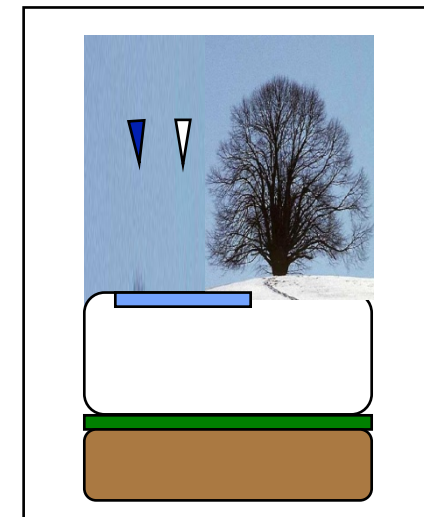
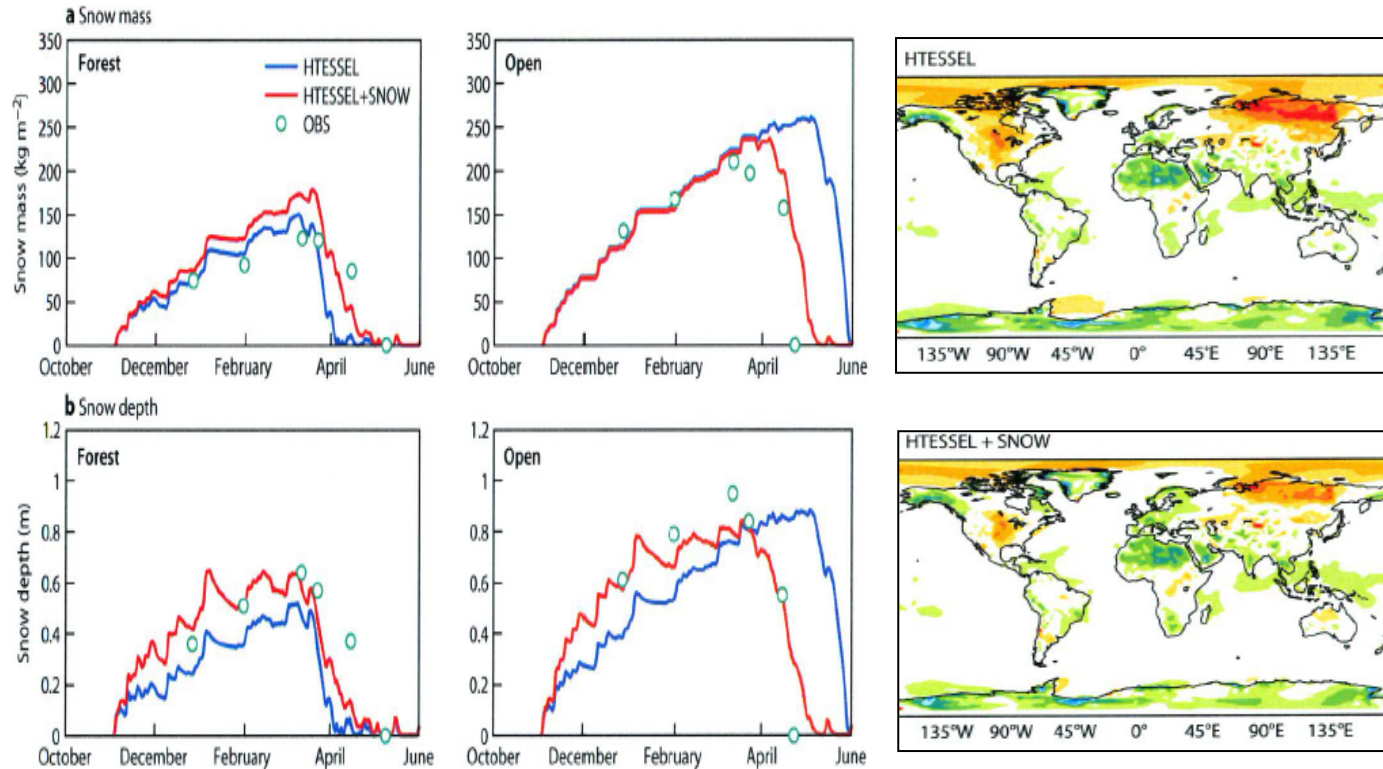
- **SL1 SNOW**

Dutra et al. (2010)

Improved snow density

Liquid water (diagnostic)

Revision of Albedo  
and sub-grid snow  
cover



The key elements of the current ECMWF snow schemes are in the treatment of snow density (including the capacity to hold liquid water content in the snowpack). The SNOWMIP 1&2 projects with their observational sites have been essential for the calibration/validation of the new scheme which was and improved with respect to the ERA-Interim snow scheme.

# A multi-layer snow scheme (research)

(Dutra et al. 2011 JGR, Dutra et al. 2012, JHM)

- **ML3 SNOW**

Dutra et al. (2012)

Up to 3 layers

Liquid water (prognostic)

Larger diurnal cycles (due to

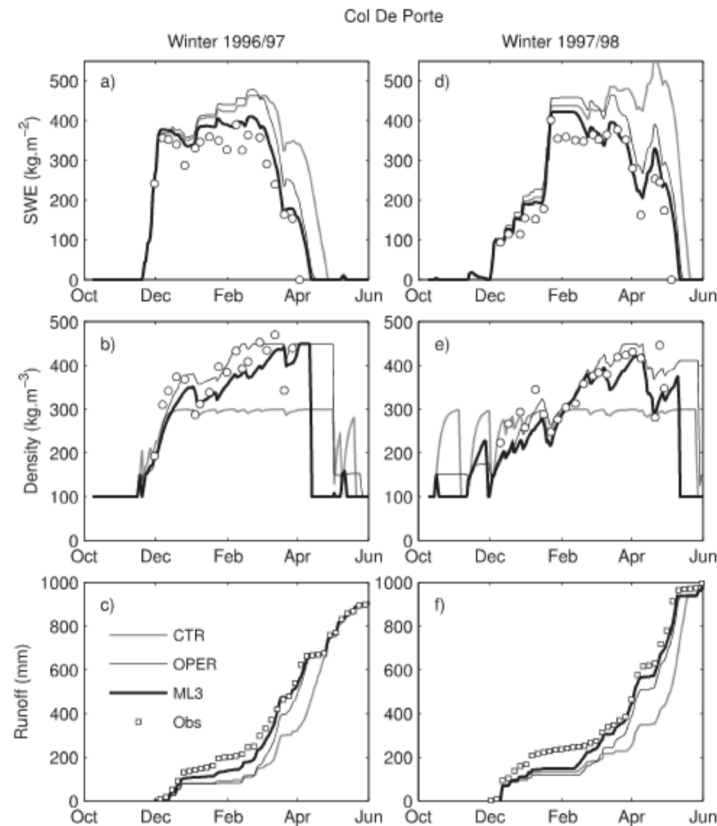


FIG. 1. Simulation results for CTR (gray), OPER (thin black), and ML3 (thick black) for the (a)–(c) 1996/97 and (d)–(f) 1997/98 winter seasons at Col de Porte site: (a),(d) snow mass, (b),(e) snow density, and (c),(f) runoff. Observations are represented by open circles. Runoff was accumulated since 1 Dec of each year and is defined as liquid precipitation and snowmelt that is in excess of the snow-cover holding capacity.

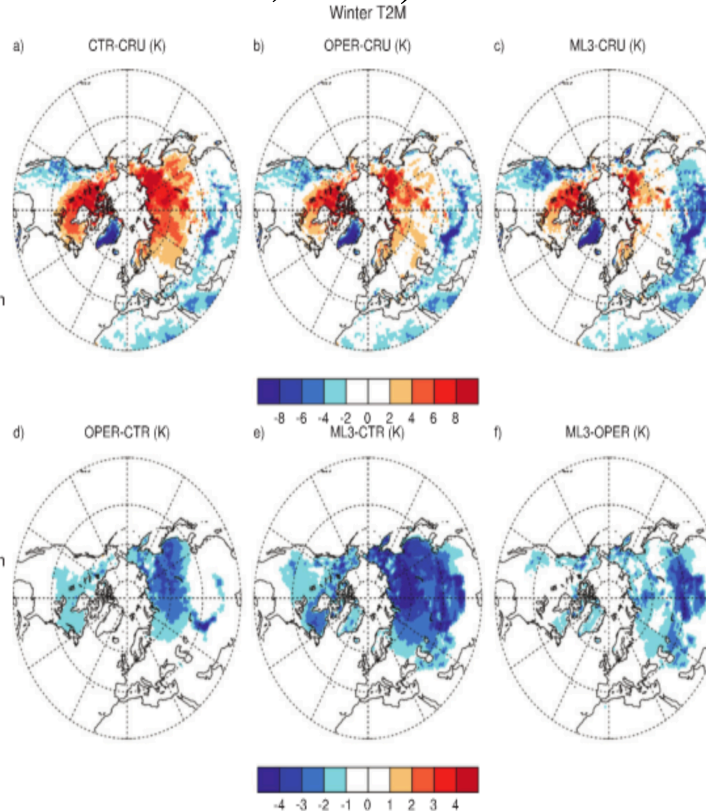
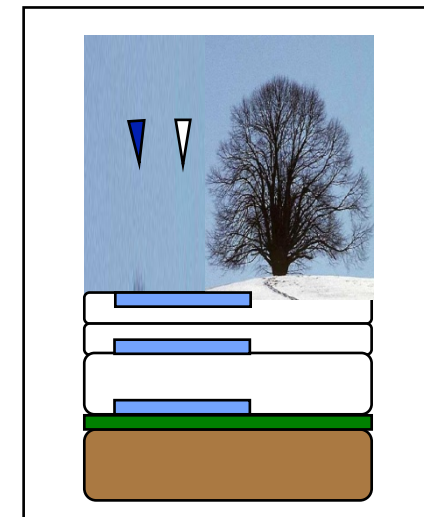
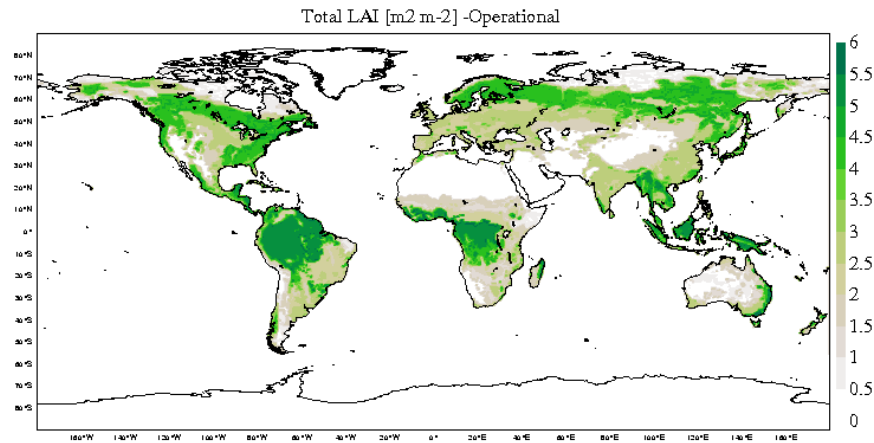


FIG. 13. Simulated winter 2-m temperature (K) biases of (a) CTR, (b) OPER, and (c) ML3 compared against CRU, and differences between (d) OPER and CTR, (e) ML3 and CTR, and (f) ML3 and OPER. Only differences significant at  $p < 0.05$  are represented. Note the different color scales between (a)–(c) and (d)–(f).



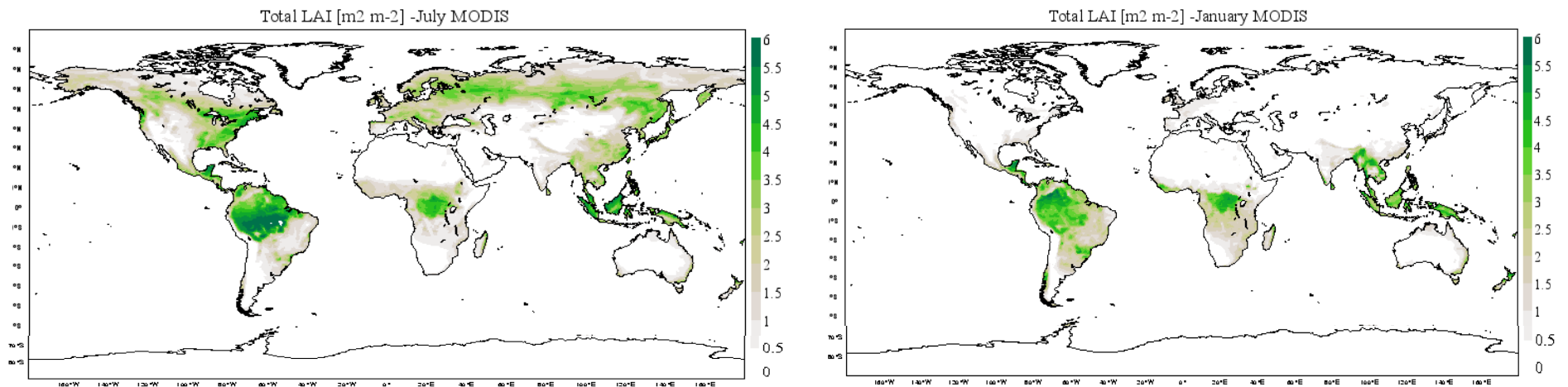
A research version multi-layer snow scheme has been developed for climate application (e.g. EC-Earth) and will be studied in Earth2Observe project. This includes up to 3 layers, an improved water cycle and further reduction of temperature bias (cooling effect in deep snow).

# Vegetation seasonality



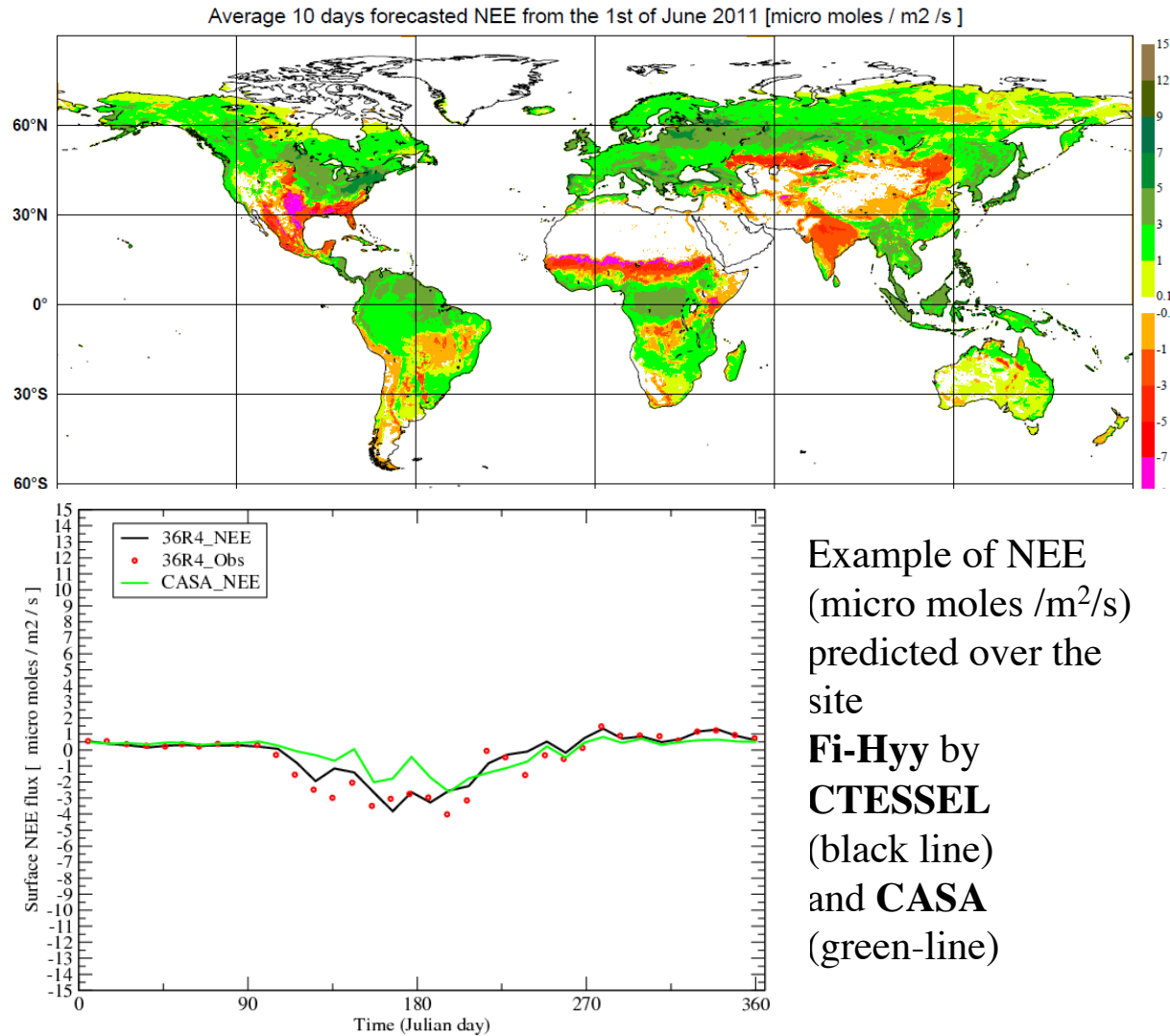
OPER LAI (van den Hurk et al. 2000, ECMWF TM)

MODIS LAI (Boussetta et al., 2011, IJRS, Myneni et al., 2002)



# Land Carbon dioxide fluxes

Boussetta et al. (2013, JGR) and ECMWF TM 675



Example of NEE (micro moles / m<sup>2</sup>/s) predicted over the site

**Fi-Hyy** by **CTESSEL** (black line) and **CASA** (green-line)

Example of Average 10 days forecast NEE (natural CO<sub>2</sub> exchange) from the 1st of June 2011 extracted from the pre operational run (e-suites) [micromoles/m<sup>2</sup>/s] – Operational from November 2011

GEOLAND-2 R&D support

## Land Natural CO<sub>2</sub>

land carbon uptake

Calvet et al. (1998)

Jarlan et al (2007)

Boussetta et al. (2013)



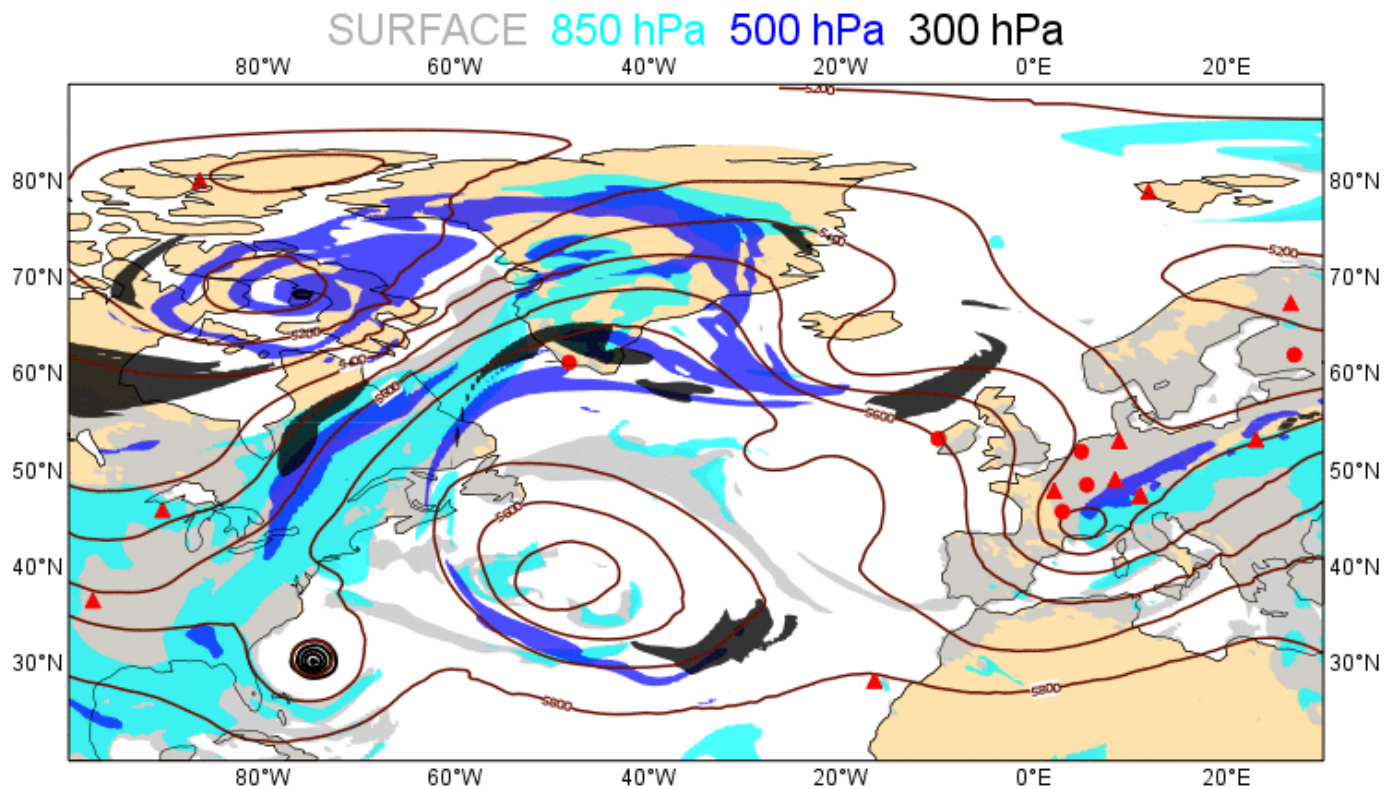
# Near Real Time CO<sub>2</sub> concentration & meteorology

Agusti-Panareda et al. (2013, ACP)

## Hurricane Sandy (2012) from different perspective

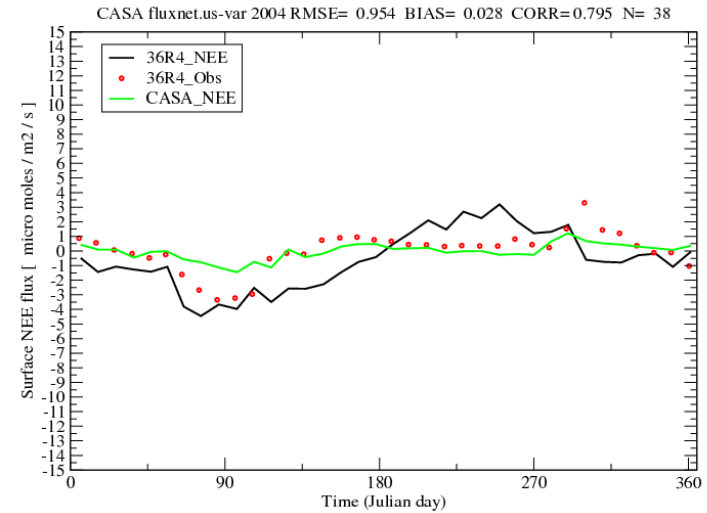
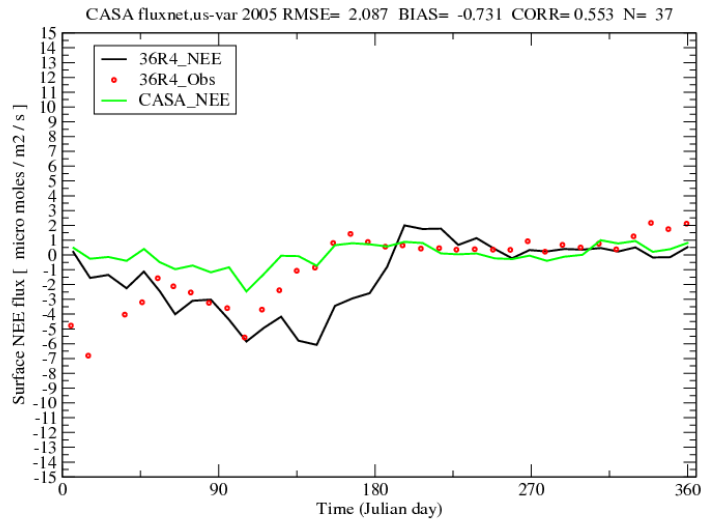
FC CO<sub>2</sub> > 379 ppm: t1279, 91 levels

2012-10-28 00:00:00



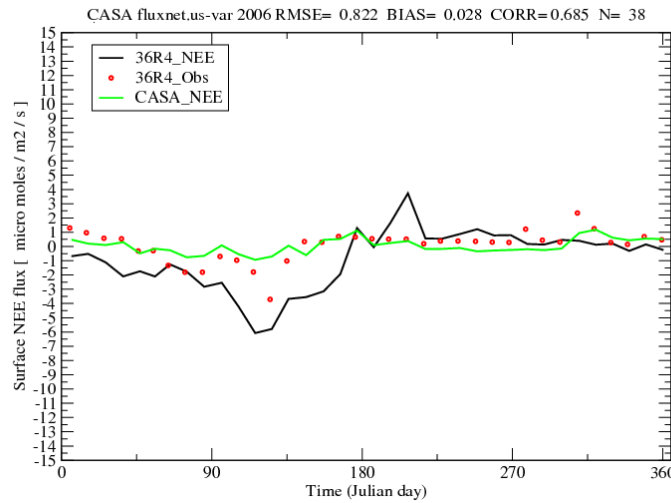


# Current shortcomings: LAI variability?



Comparison of 3 different years show the current shortcoming of the use of LAI climatology => presence of large interannual variability? Harvest period not matched?

NEE (micro moles /m<sup>2</sup>/s) predicted over the site US-Var by **CTESSEL** (black line) and **CASA** (green-line)



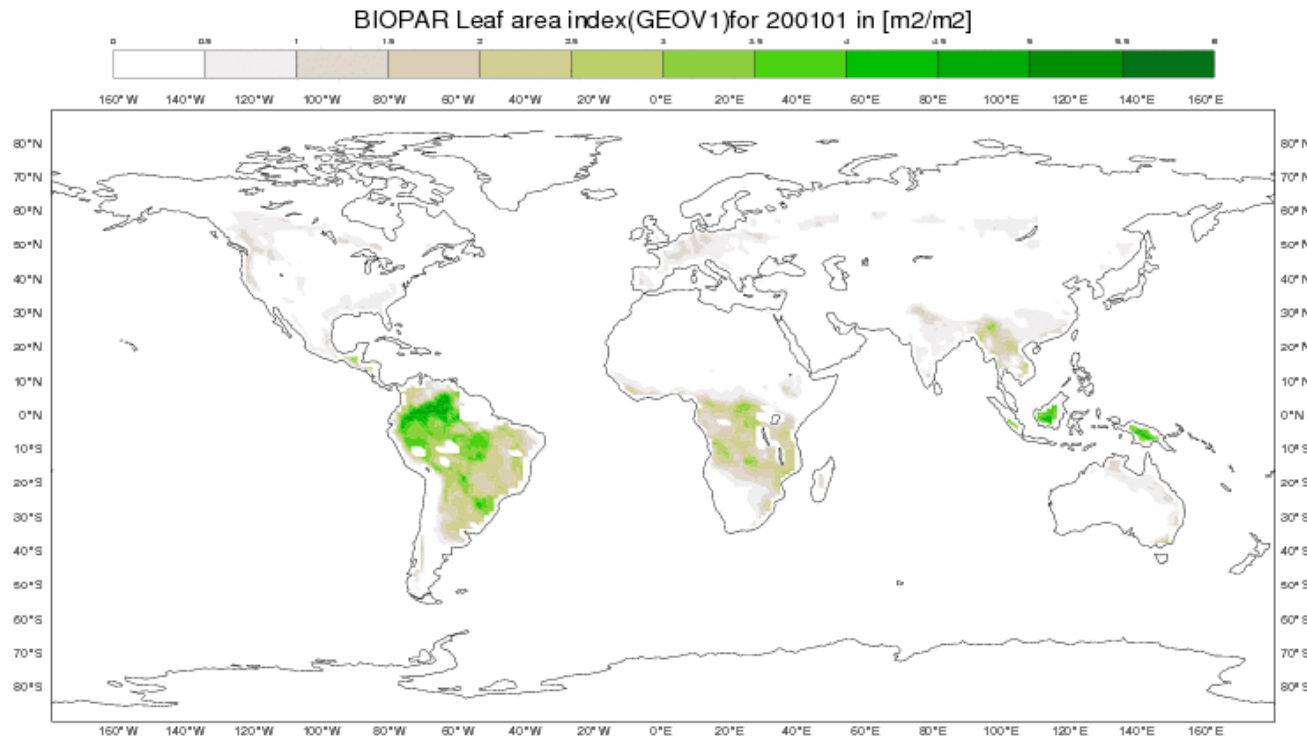
# Vegetation inter-annual variability (R&D)

geoland:2

Imagine



A 10-year global LAI (and Albedo) product (10-day frequency) has been provided by the geoland2 BIOPAR (GEOV1 product) and continues in IMAGINES project. Our plan is to assimilate in the ECMWF system to test the impact fluxes and forecasts. Preparing for SENTINEL ESA satellites.



<http://fp7-imagines.eu/>



# Summary & Outlook

- **The ECMWF land surface scheme and its evolution**

- Uses the tiling concept to represent sub-grid land variability including forest and forest+snow tiles, and upcoming a lake dedicated tile (cy40r1)

- **Benefits of the tiling in presence of large contrasts**

- Each tile has its process description (no ad-hoc or effective parameters)

- **Shortcomings of the tiling**

- No surface boundary layer mixing (blending height hypothesis)
- Too strong decoupling of snow surface (2m temperature forest bias)
- Single soil layer underneath

- **Outlook: How approach calls for complexity? What comes first?**

The enhanced representation of snow and vegetation + introduction of a SBL scheme to account for mixing are foreseeable developments with NWP relevance

## References

PDFs of published articles available as ECMWF TM or via web-sharing networks

# Land surface ongoing & future developments

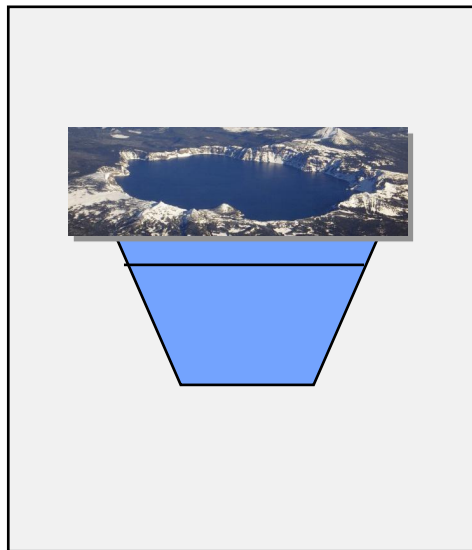
2013	2014-2015	2015-2020
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- **FLake**

Mironov et al (2010),  
 Dutra et al. (2010),  
 Balsamo et al. (2010)  
 Balsamo et al. (2011)

Extra tile (9) to account  
 for sub-grid lakes

Lake Climatology used in  
 S4



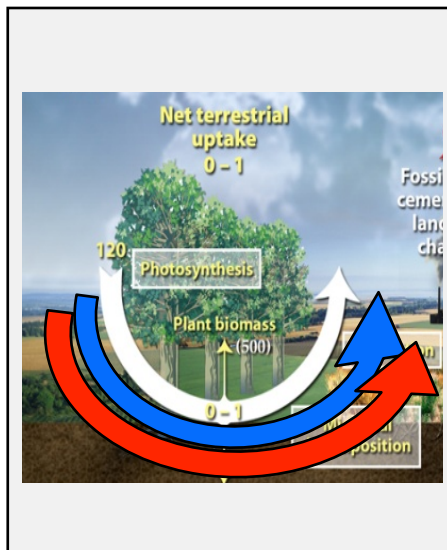
- **H<sub>2</sub>O / E / CO<sub>2</sub>**

Carbon-driven vegetation

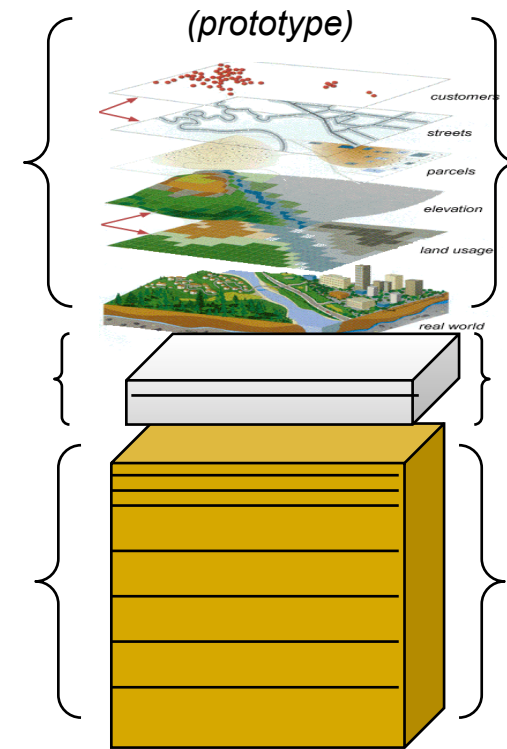
scheme at the surface,

Atmospheric CO<sub>2</sub> + DA

(FP7 & Copernicus)



- Towards **Interactive Ecosystem modelling** to respond to several applications needs



# Land surface data assimilation status

1999	2004	2010/2011
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## OI screen level analysis

Douville et al. (2000)  
Mahfouf et al. (2000)

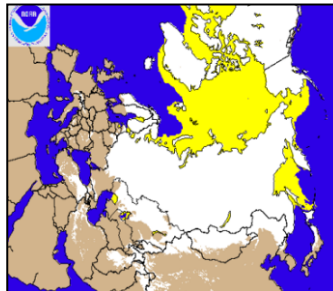
**Soil moisture 1D OI analysis**  
based on Temperature and  
relative humidity analysis



SYNOP Data

## Revised snow analysis

Drusch et al. (2004)  
Cressman snow depth  
analysis using SYNOP data  
improved by using NOAA /  
NSEDIS Snow cover extend  
data (24km)



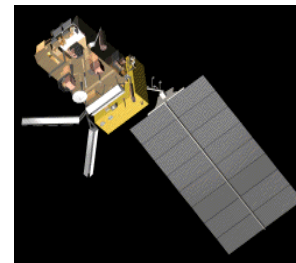
NOAA/NESDIS IMS

**Optimum Interpolation (OI) snow analysis**  
**Pre-processing NESDIS data**  
**High resolution NESDIS data (4km)**  
de Rosnay et al., 2012

## SEKF Soil Moisture analysis

Simplified Extended Kalman Filter  
Drusch et al. GRL (2009)  
de Rosnay et al (2012)

## Use of satellite data



**METOP-ASCAT**

de Rosnay et al., 2011



**SMOS**

Sabater et al., 2011

## Validation activities

Albergel et al. 2011, 2012, 2013

# The land water storage verification

The soil moisture and the snow depth produced with the current land surface model simulation are part of the new “ERA-Interim/Land” dataset available from ECMWF.

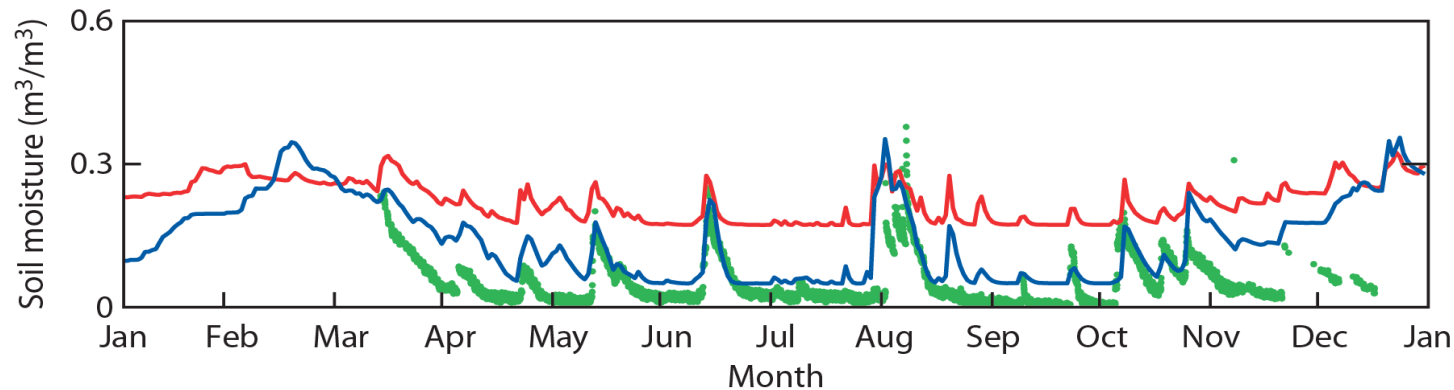


Figure 1: Evolution of volumetric soil moisture at a site in Utah for the year 2010. In-situ observations in green, ERA-Interim estimates in red, and ERA-Interim/Land estimates in blue.

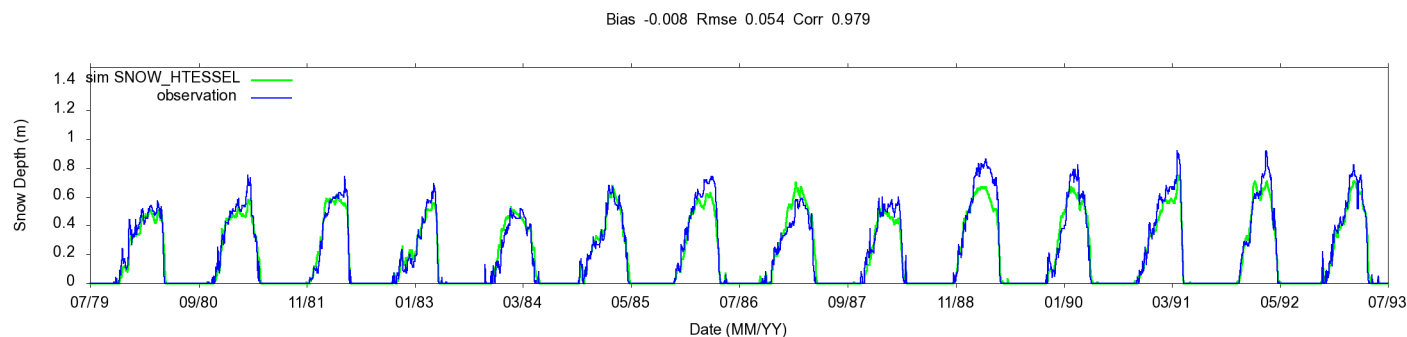


Figure 2: Long term evolution of the HTESSEL snow-depth (in green) compared with in-situ measurements (in blue) from 1979 to 1993 at Perm station (58.0N, 56.5E).

# The land water fluxes verification

The ERA-Interim/Land fluxes are validated with independent datasets used as benchmarking.

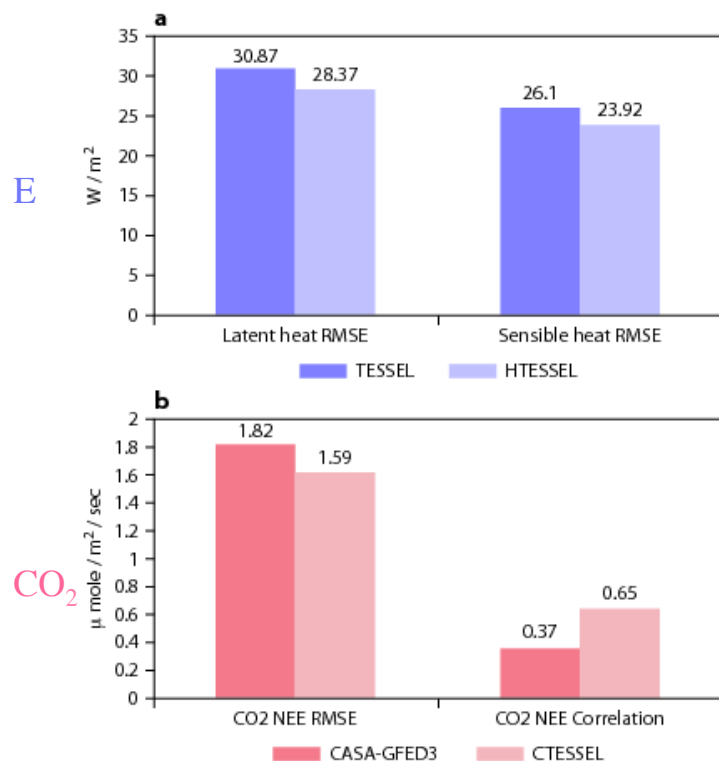


Figure 1: Mean errors over 36 stations with hourly data from FLUXNET & CEOP networks (decadal-averages) in 2006 for HTESSSEL, TESSEL scheme versions (with latent and sensible heat fluxes RMSE shown in (a)) and HTESSSEL and CHTESSSEL (with Net Ecosystem Exchange RMSE and Correlation in (b)).

## Validation of $H_2O$ / E / $CO_2$ cycles

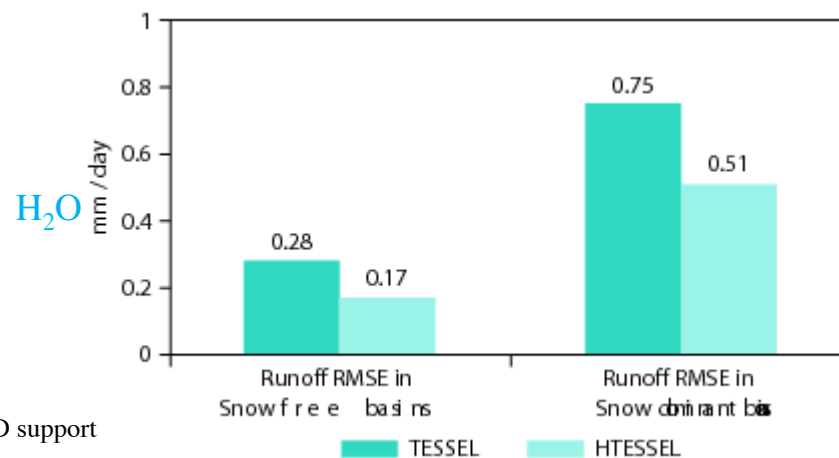
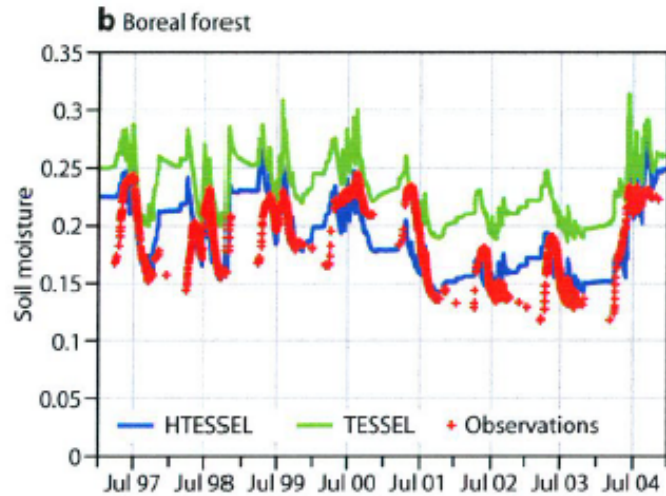


Figure 2: Runoff root-mean-square error (RMSE) for GSWP2 global offline land simulations (1986–1995) verified with GRDC monthly river discharge observations on mainly snow-free basins (North-East and Central Europe) and snow-dominated basins (Yukon, Podka, Lena, Tom, Ob, Yenisei, Mackenzie, Volga, Irtish and Neva). The mean RMSEs are area-weighted and show the TESSEL and HTESSSEL scheme versions.

GEOLAND-2 R&D support

# A revised soil hydrology

(Balsamo et al. 2009, JHM)



Long record observations at BERMS-Canadian site have crucial to assess the hydrological performance of the new scheme

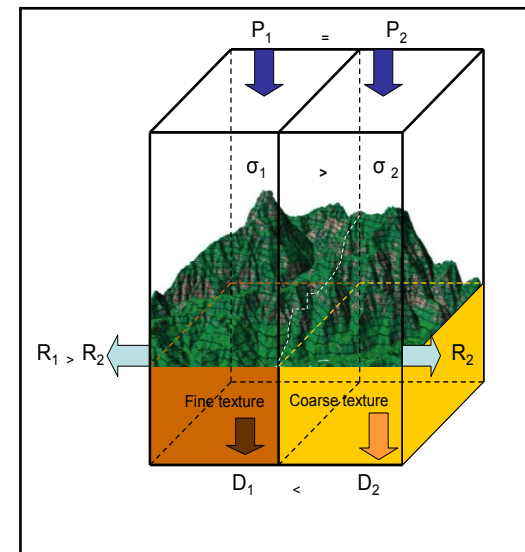
- Hydrology-**TESSEL**

Balsamo et al. (2009)  
van den Hurk and Viterbo (2003)

Global Soil Texture (FAO)

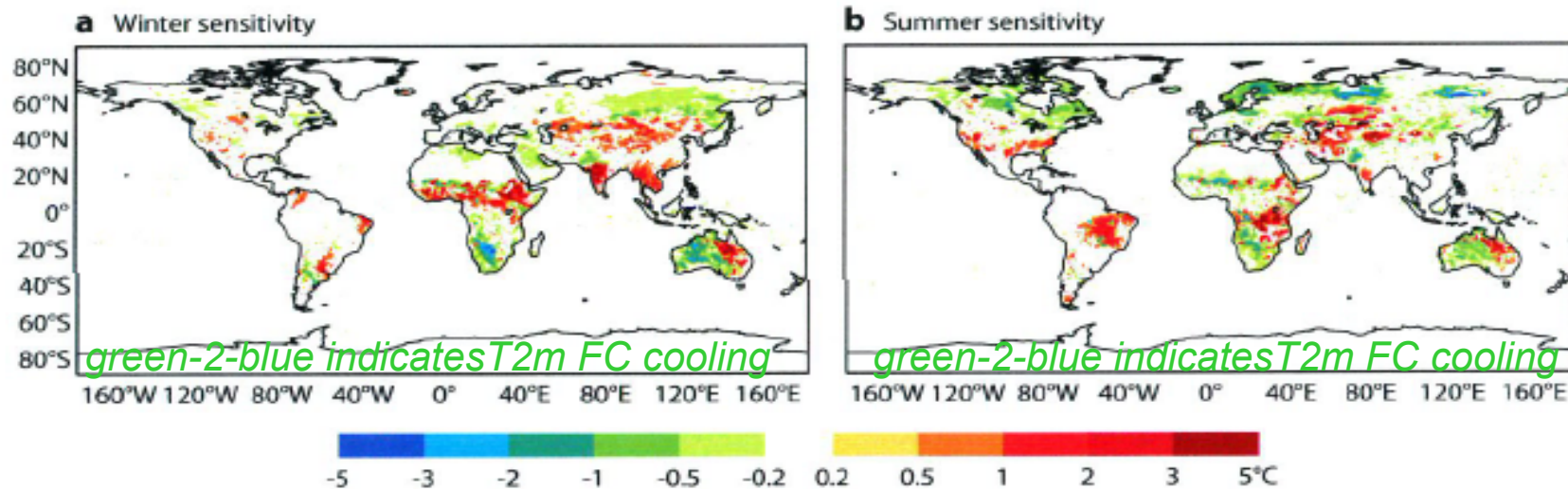
Van Genuchten hydraulic properties

Variable Infiltration capacity & surface runoff revision

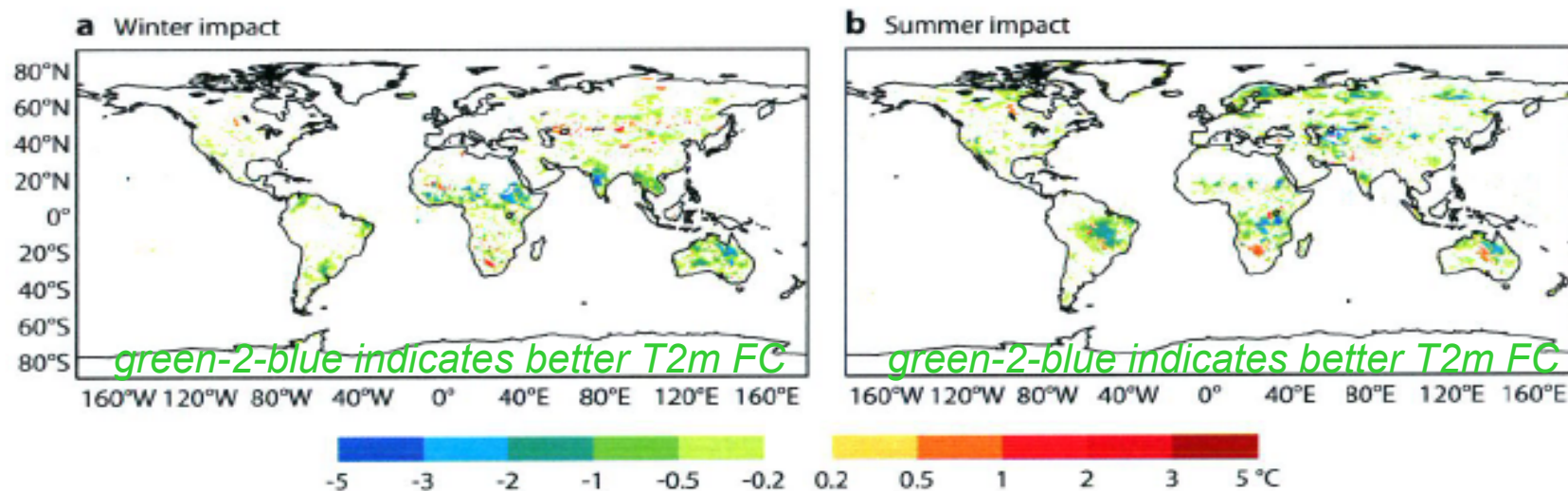




# Forecasts sensitivity and impact

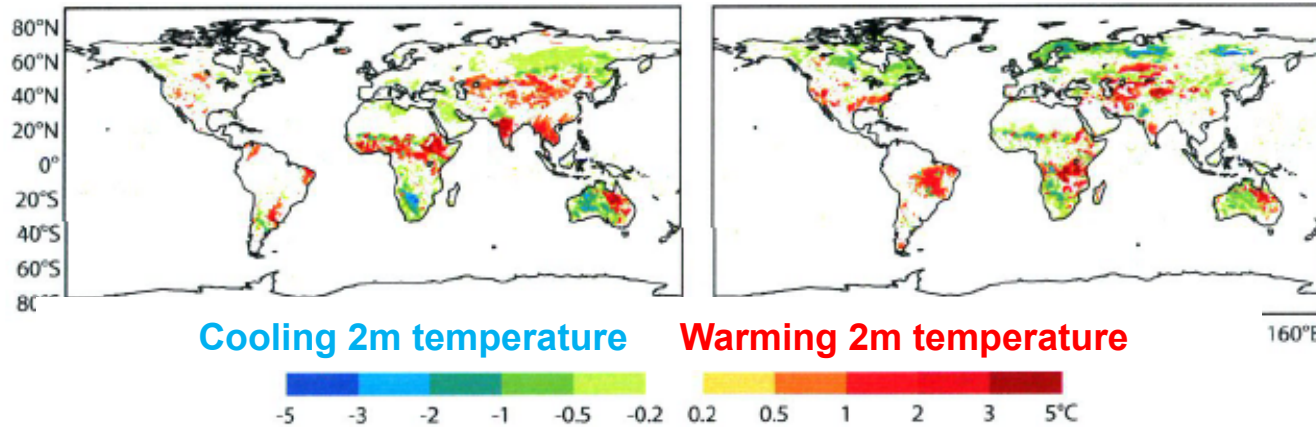


- The revised soil/snow scheme introduce additive improvements respectively in summer/winter seasons forecasts of 2m temperatures

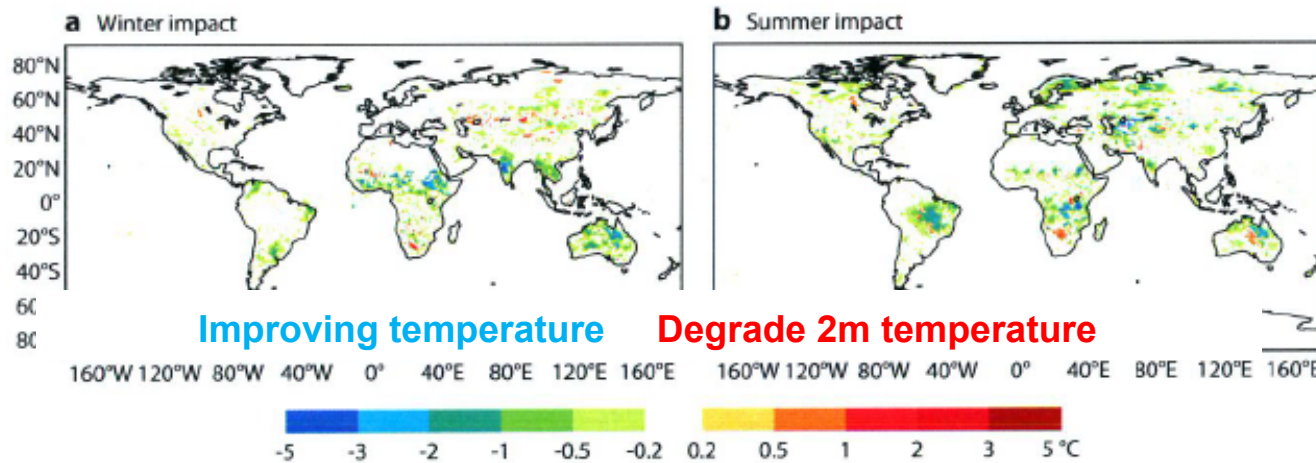


# Forecasts sensitivity and impact to land

Sensitivity of a set of T2m Day-2 forecasts in winter 2008 (DJF) and Summer 2008 (JJA)



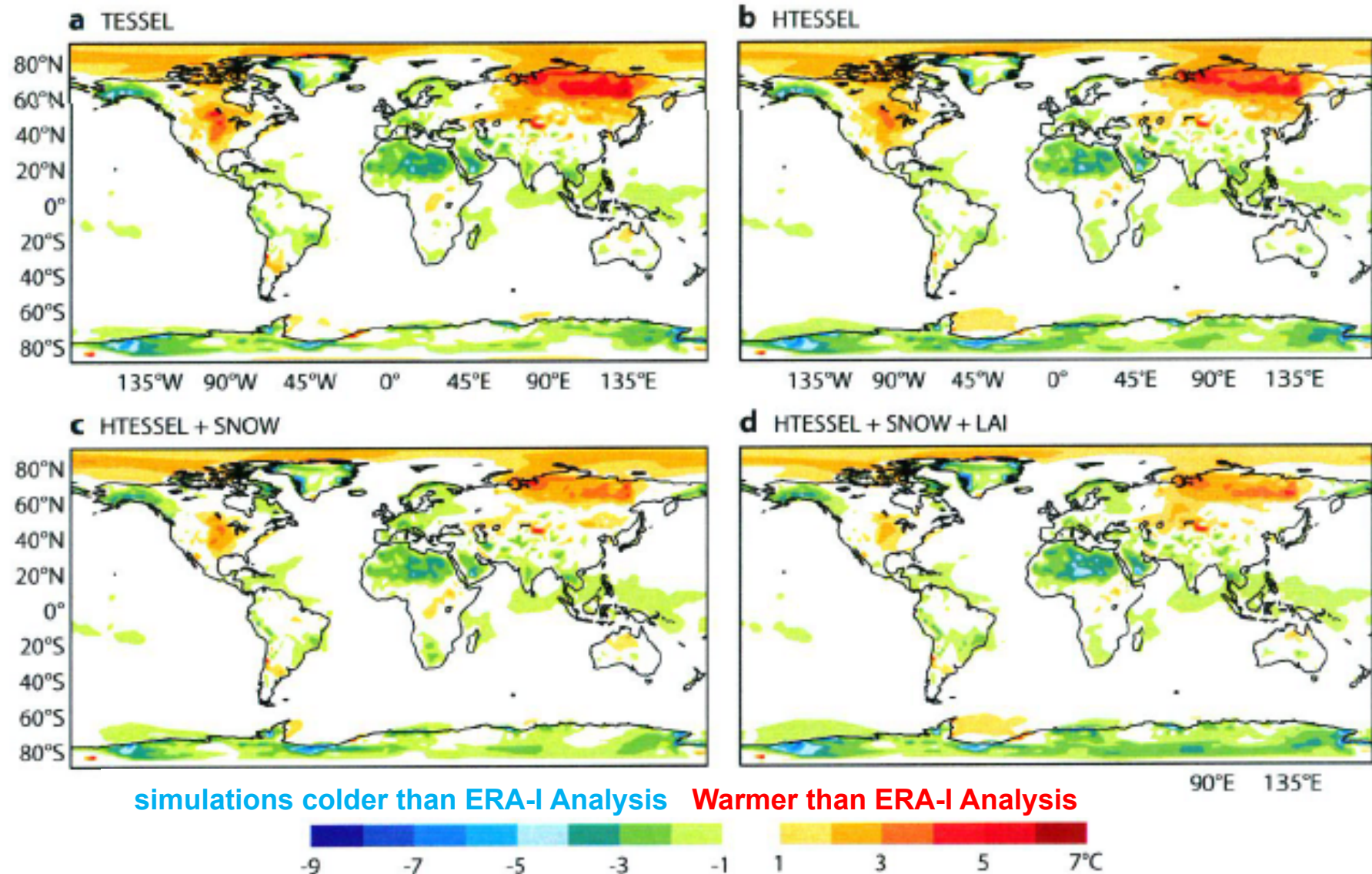
Forecast Impact (Mean Absolute Error reduction of the T2m Day-2 forecast error)



The revised IACM faces workshop, 36R4 is compared to the land surface model version (CY31R2 LSM used in ERA-Interim) for its sensitivity and impact on the short-term weather forecasts of 2m temperature showing an improvement also in Day-2 range

# Land-related improvements in climate runs

Hindcast (13-months integrations with specified daily SSTs). Here shown the evolution of the annual mean T2m errors compared to analysis



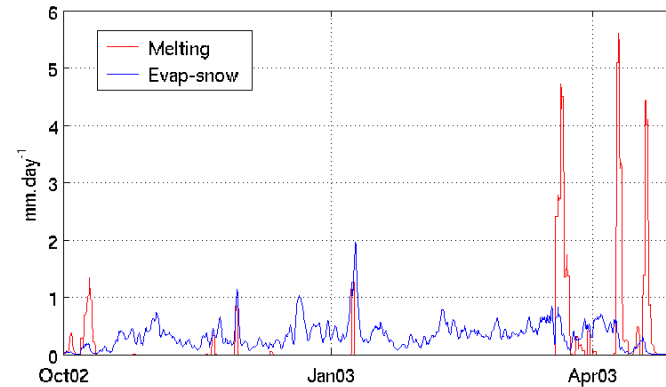
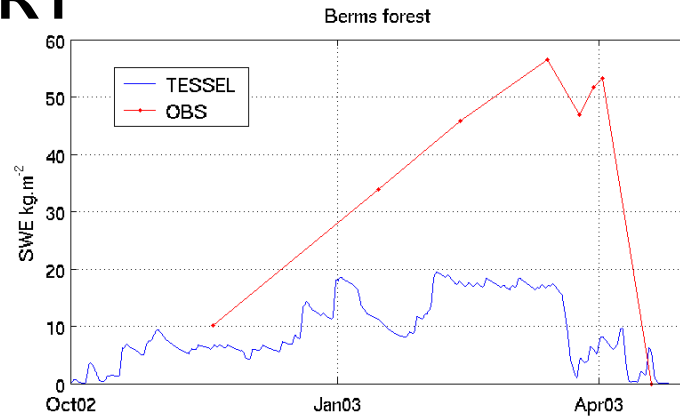
The revised land surface scheme in CY30R4 (d) is compared to the land surface model version (a, CY31R2 LSM used in ERA-Interim) for its impact on long-range forecasts of 2m temperature showing an improvement on annual mean 2m temperature

# Impact of Forest+Snow Roughness Length

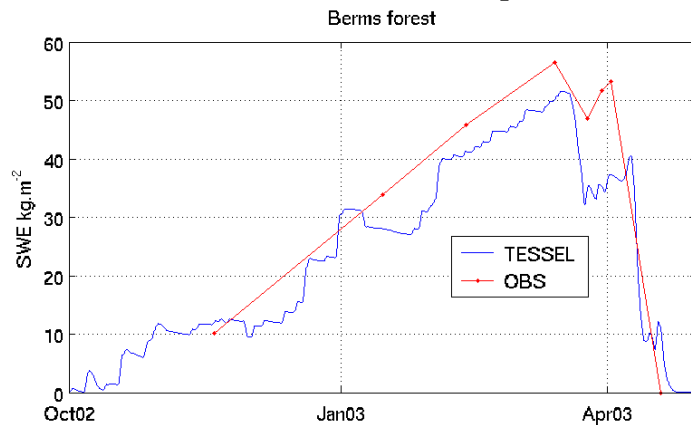
Dutra et al. 2009

The introduction of a vegetation dependent roughness length affecting the aerodynamic resistance show sensitivity on snow accumulation (less sublimation)

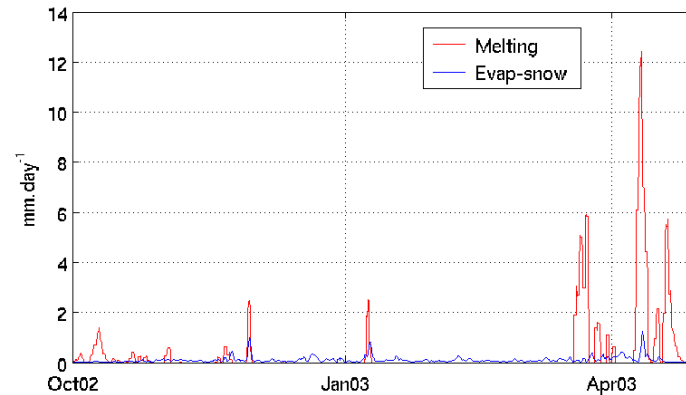
### 30R1



### 31R2, snow water equivalent



### melting & sublimation



## SnowMIP2: BERMS forest site simulations

COLA-GMU LSM Workshop, 6 Dec 2013, G. Balsamo

