

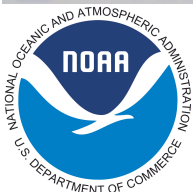
Land-Atmosphere Interaction *(and Cloud Formation)*

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*COLA Workshop on Land Surface Modeling in Support of
NWP and Sub-Seasonal Climate Prediction
5-6 December 2013, George Mason University, Fairfax, VA*



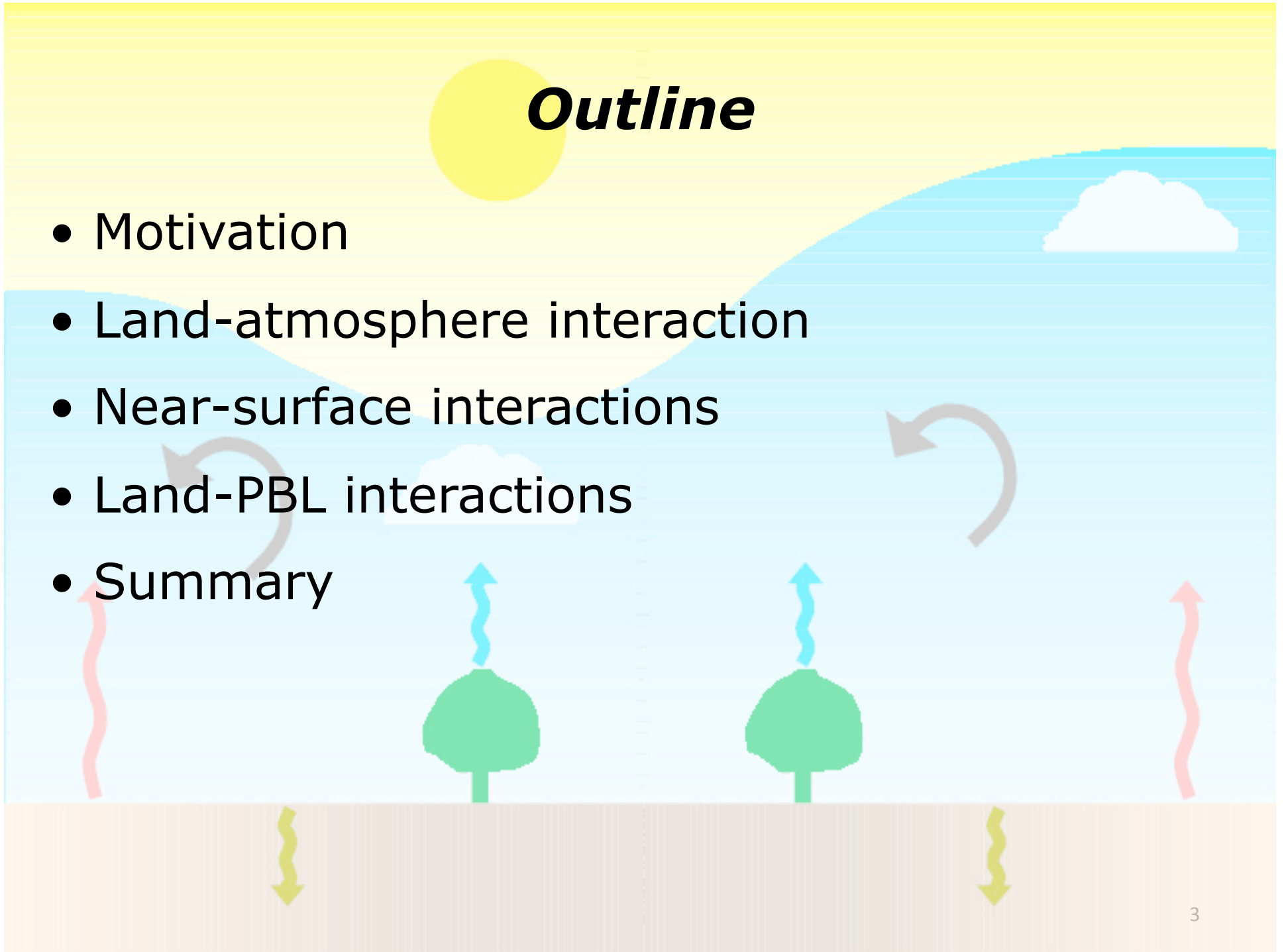
NOAA Center for Weather and Climate Prediction (NCWCP), College Park, Maryland, USA



World Weather Building

Outline

- Motivation
- Land-atmosphere interaction
- Near-surface interactions
- Land-PBL interactions
- Summary



Motivation

- **Land-atmosphere interaction** and coupling strength remain weak links in current land-surface and atmospheric prediction models.
- Coupling strength affects surface fluxes, so important for weather and climate.
- We need to understand the many land and atmospheric processes and interactions, with proper representation in weather and climate models.
- ***Coupling begins locally.***

Land-Atmosphere Interaction

*Betts
(1996)*

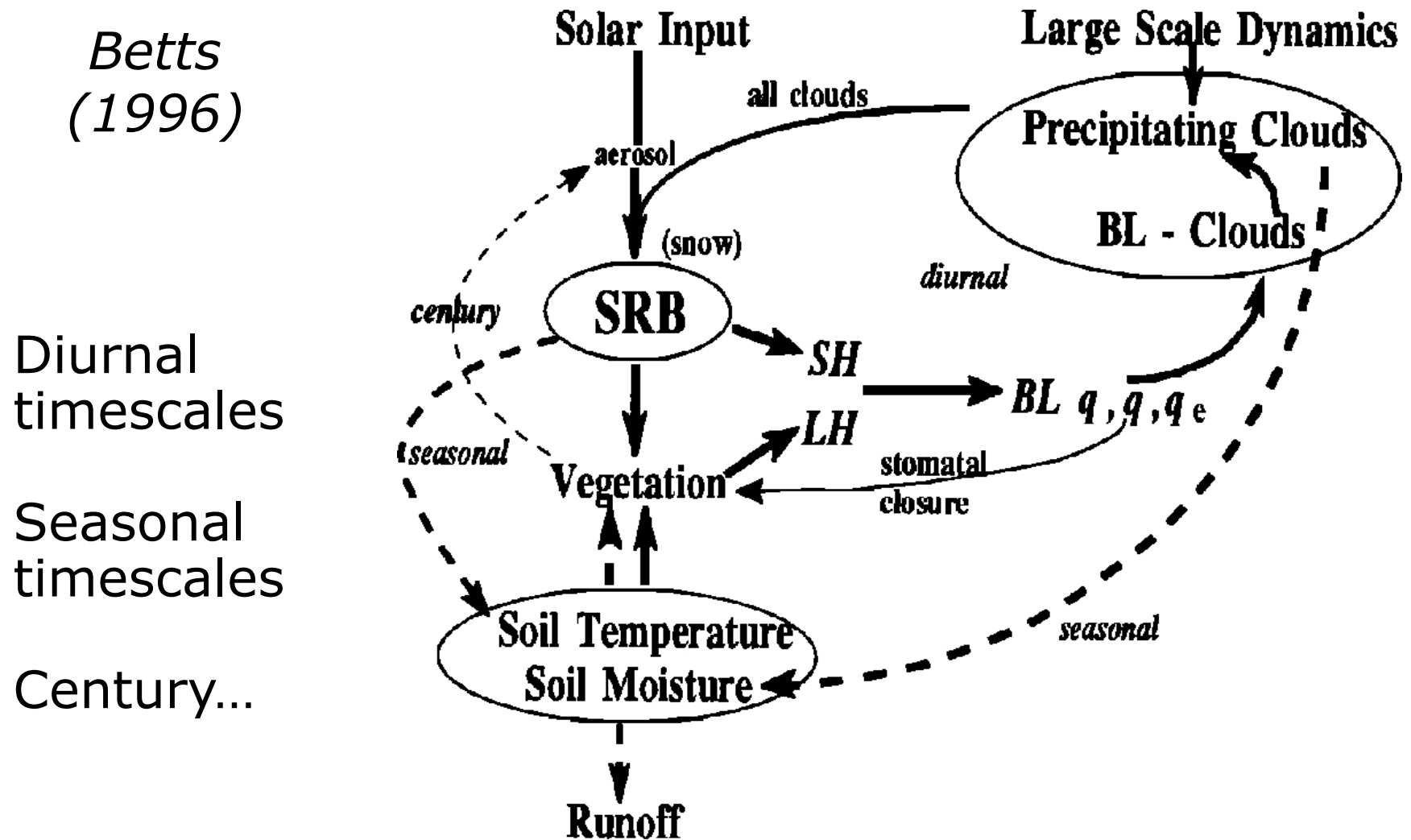


Figure 1. Schematic showing some important land surface-atmosphere interactions on different timescales.

Land-Atmosphere Interaction

Beljaars
(2005)

EC model/
TESSEL

“We discussed including this in a recent document, but dropped it because it was too confusing.”

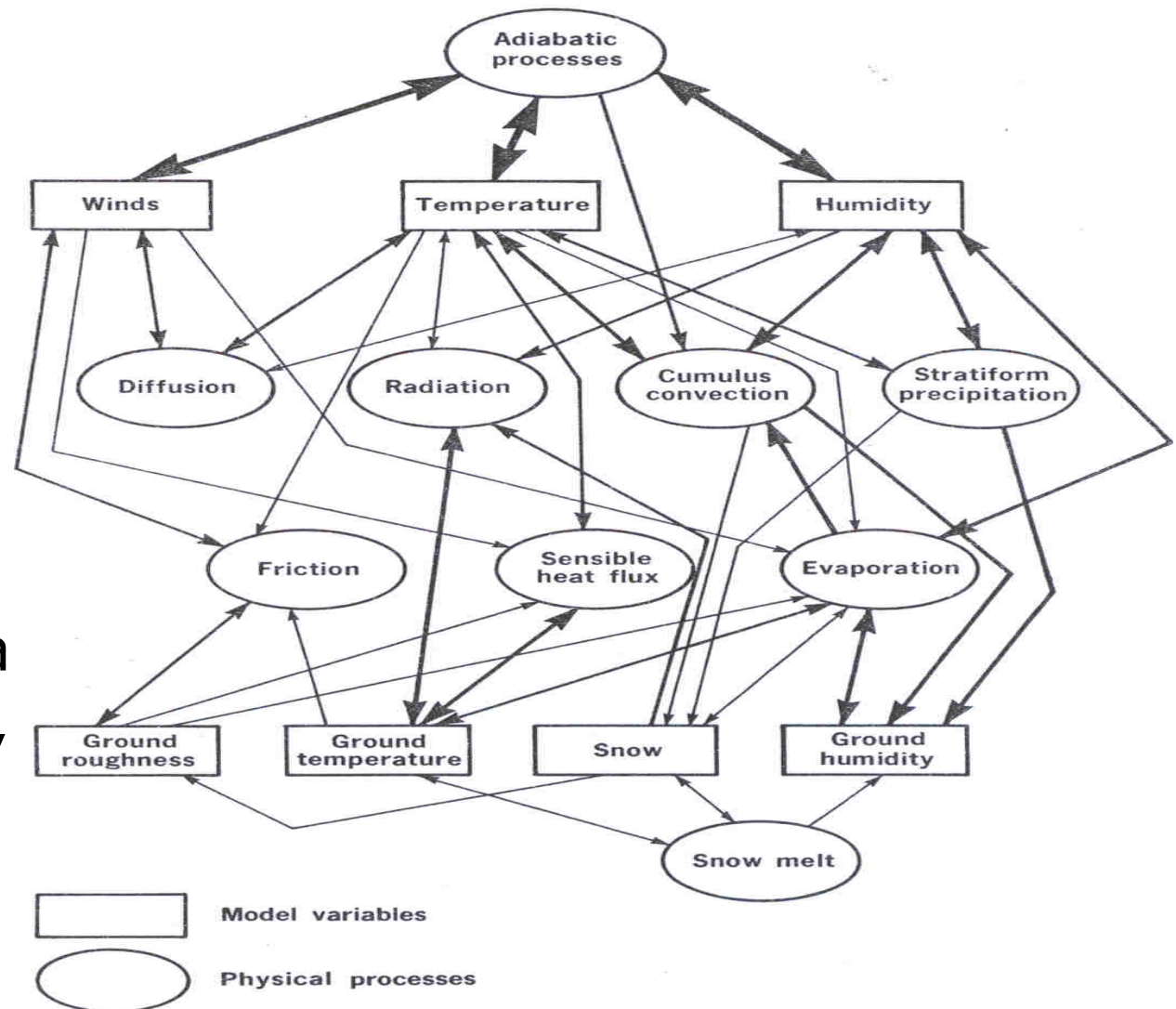
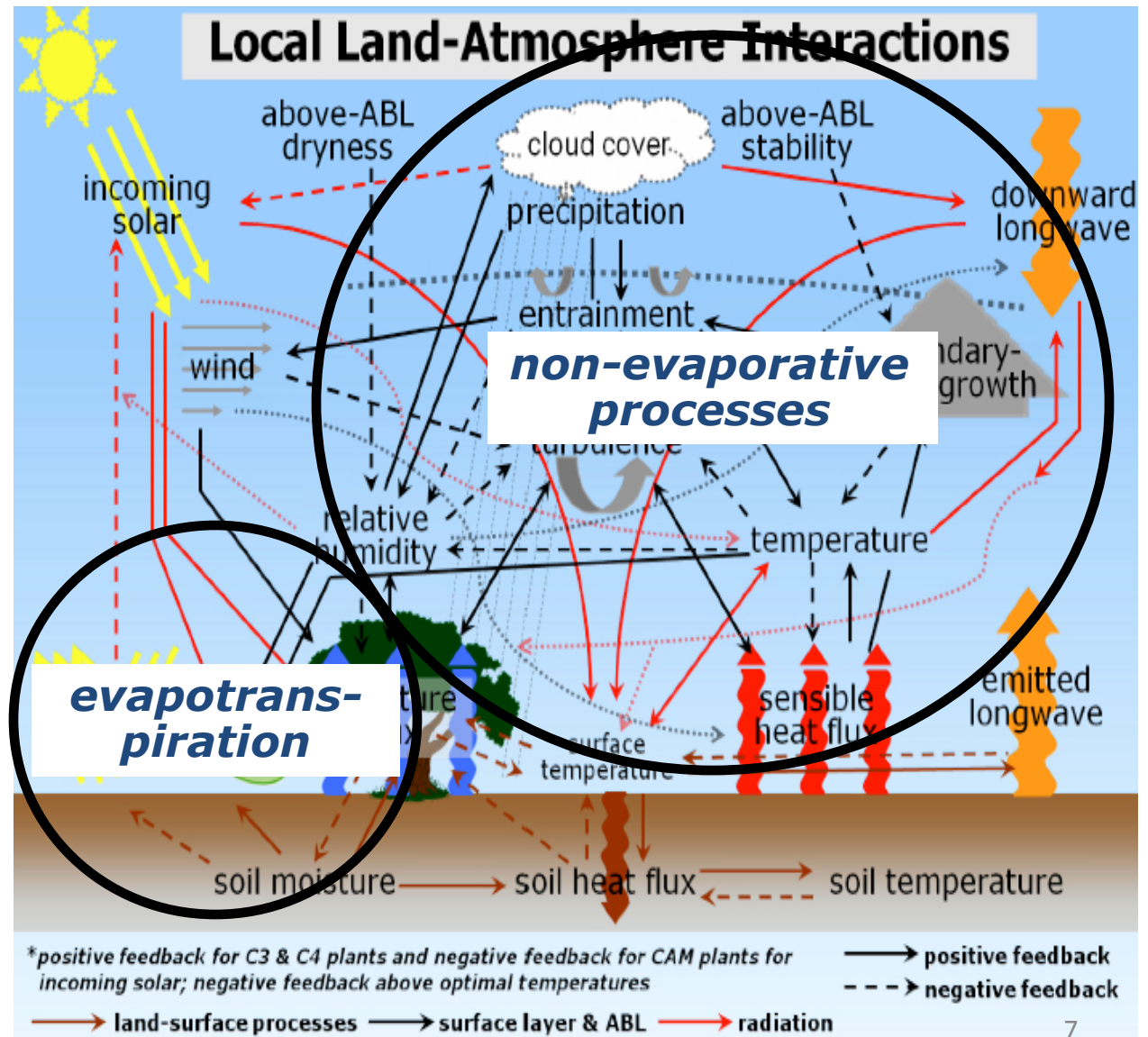


Fig. 1.1 Schematic representation of the processes included in the ECMWF model.

Land-Atmosphere Interaction

Adapted from
Ek & Holtslag
(2004)

Characterized many land and atmospheric processes and feedbacks for typical daytime with focus on soil moisture vs other processes.



"GEWEX Imperatives:
Plans for 2013 and
Beyond" (gewex.org)

Land-Atmosphere Interaction

*van Heerwaarden et al
(2009)*

Negative feedback mechanisms and the relationships among variables that regulate evaporation.

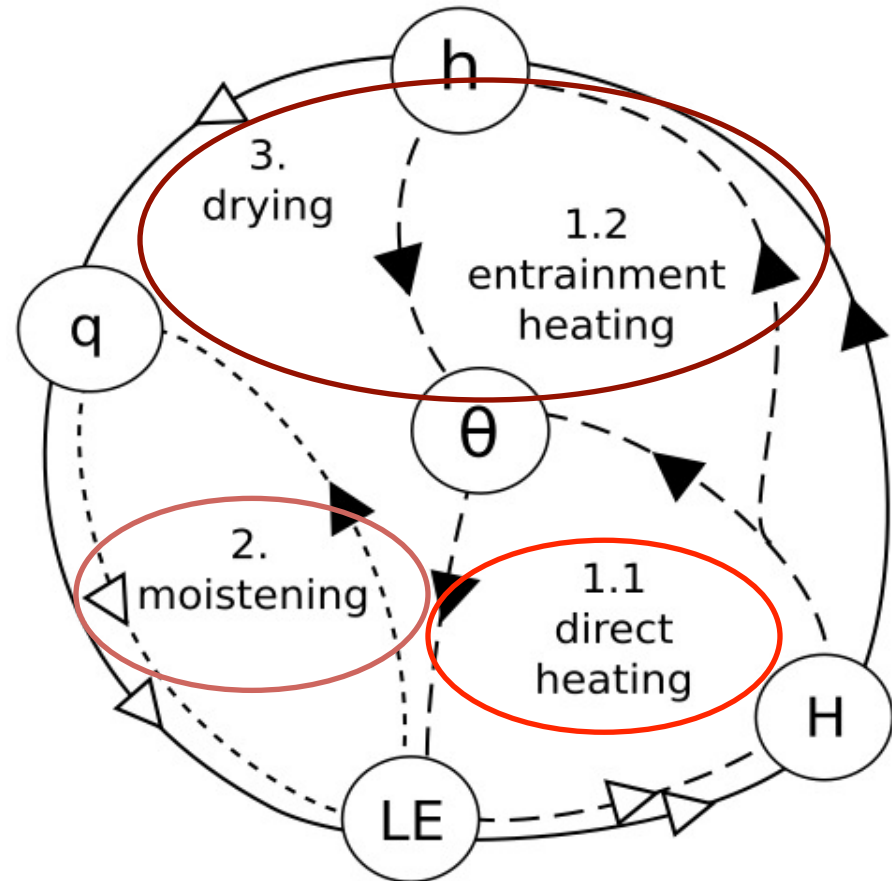
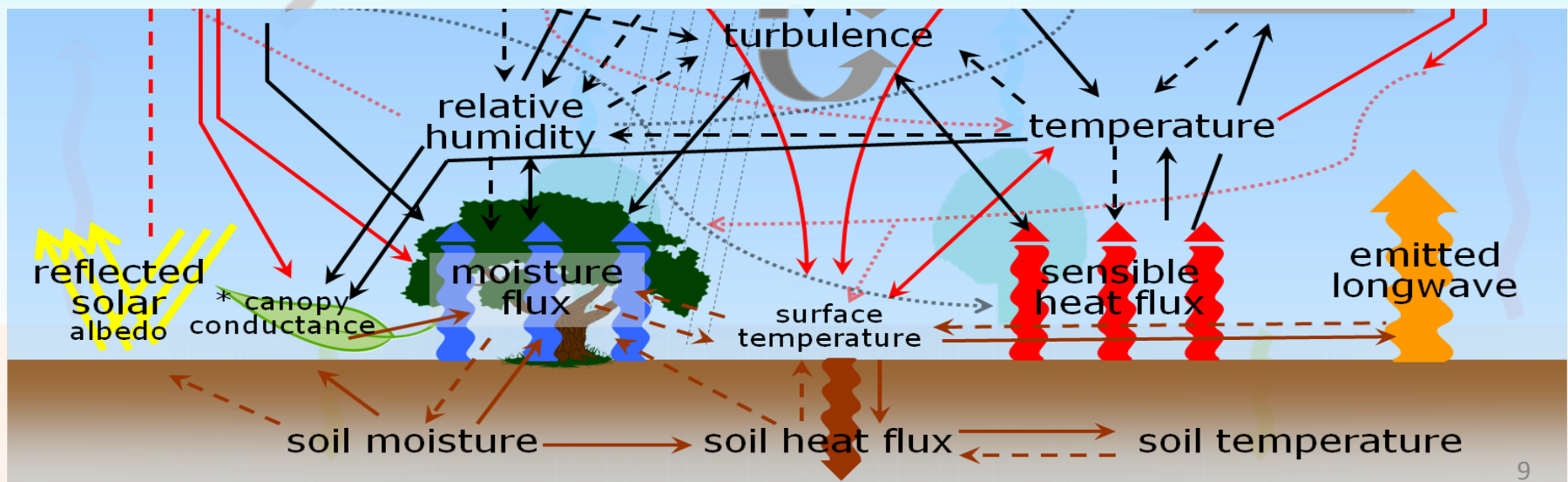


Figure 1. Feedbacks in the coupled land-atmosphere system. Closed arrows represent positive relationships, and open arrows negative relationships. Each of the three feedbacks has a distinct line style. LE is the latent heat flux, H is the sensible heat flux, θ is the bulk potential temperature of the CBL, q is the bulk specific humidity of the CBL and h is the CBL height.

Near-Surface Interactions

- What is nature of **near-surface** land-atmosphere coupling? For strong (weak) coupling, a given soil moisture change yields large (small) ET change.
- What is relationship between soil moisture and ET (or ef) in terms of near-surface turbulence, atmospheric variables, vegetation and soil processes.
- Expand work of Jacobs et al (2008), Jarvis (1985) et al, and others.



Near-Surface Interactions:

Soil moisture – transpiration relationship

Evaporative fraction for transpiration:

$$ef_t = \frac{s + \frac{\rho c_p g_a \delta e}{R_n - G}}{s + \gamma \left(1 + \frac{g_a}{g_c}\right)},$$

Penman-Monteith

Evap. fraction change with soil moisture change:

$$\frac{\partial \ln ef_t}{\partial \Theta} = \frac{1}{\delta \Theta_{rz}} \left\{ \left[\left(\frac{s + \gamma}{\gamma} \right) \frac{g_c}{g_a} + 1 \right]^{-1} + \left[\frac{s(R_n - G)}{\rho c_p g_a \delta e} + 1 \right]^{-1} \frac{\delta \Theta_{rz}}{\Theta_{ns}} \frac{b \beta G}{(R_n - G)} \right\}$$

ga/gc-term: stomatal control vs sfc-layer turbulence, range: 0-1 (J08, J85)

vG-term: soil heat flux contribution, 0 to 0(1)

Stronger Coupling:

Strong stomatal control, strong turbulence, e.g. forest with dry soil

Strong turbulence, dry air, large G, small Rn, large soil heat flux, wet soil

Weaker Coupling:

Weak stomatal control, weak turbulence, e.g. grassland with wet soil

Weak turbulence, moist air, small G, large Rn, small soil heat flux, dry soil

Near-surface Interactions:

Vegetated

**Stronger
Coupling**

“ ωv ” coupling parameter
= $g_a/g_c + vG$ terms

Additional factors:

- $T \uparrow$ $RH \downarrow$ $R_n \downarrow \rightarrow \omega v \uparrow$
- forest > grass
- sand > clay
- g_a/g_c term \gg vG term (generally)

**Fluxnet Data Sets
with good soil moisture
measurements**

**Weaker
Coupling**

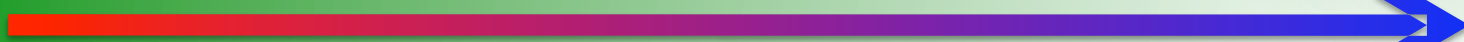
strong



TURBULENCE

weak

dry

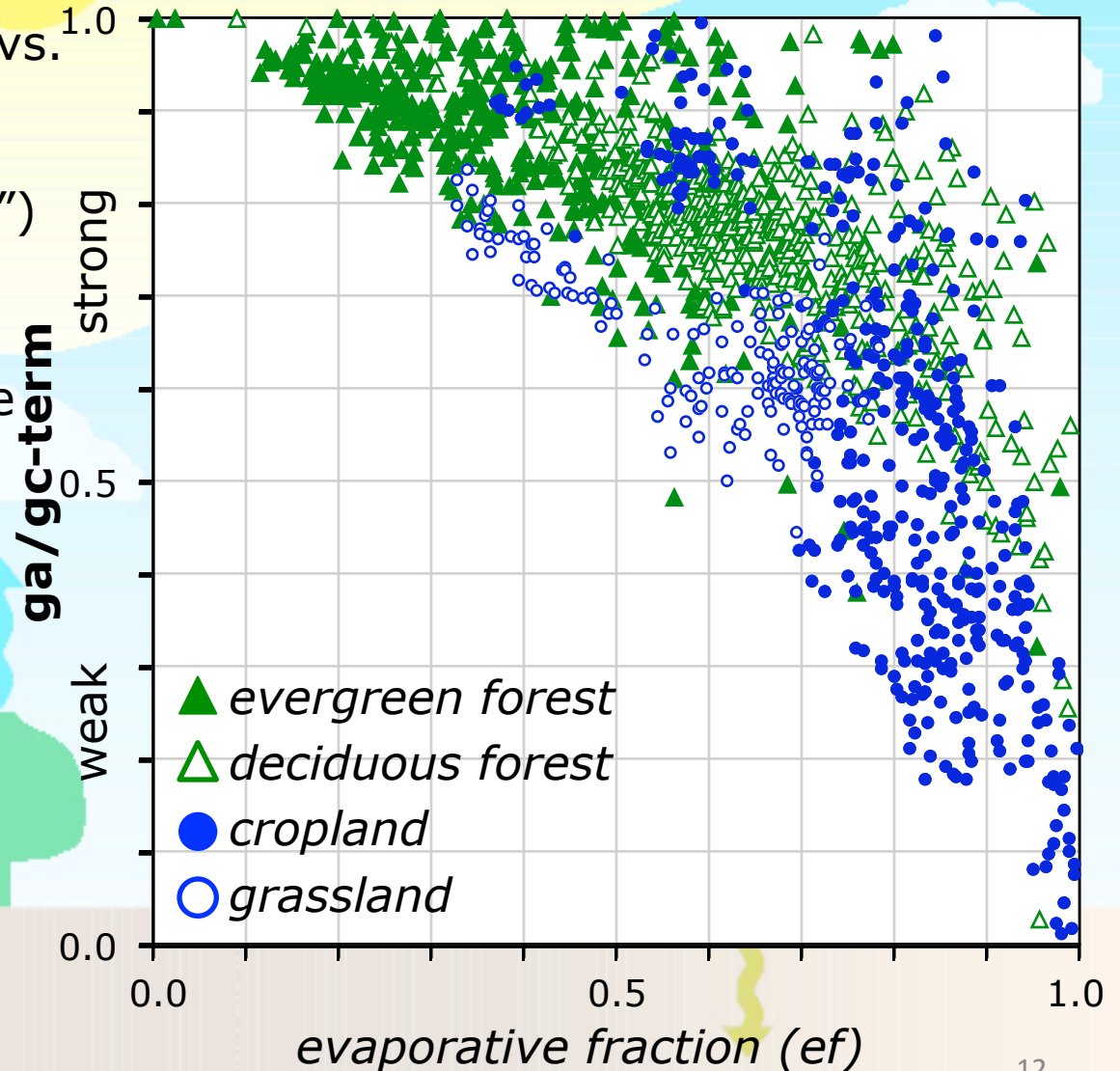


SOIL MOISTURE

wet

Near-Surface Interactions: Evaporative Fraction vs Soil Moisture

- evaporative fraction (ef) vs. **g_a/g_c -term** (coupling strength) from surface flux site observations ("Fluxnet")
- higher ef:
 - stronger land-atmosphere coupling for forests.
 - weaker land-atmosphere coupling for cropland and grassland.
- lower ef: strong coupling regardless of vegetation type: due to stronger surface heating and turbulence (larger g_a , smaller g_c).



Need to include vG-term

Near-Surface Interactions:

Soil moisture – “bare” soil evap relationship

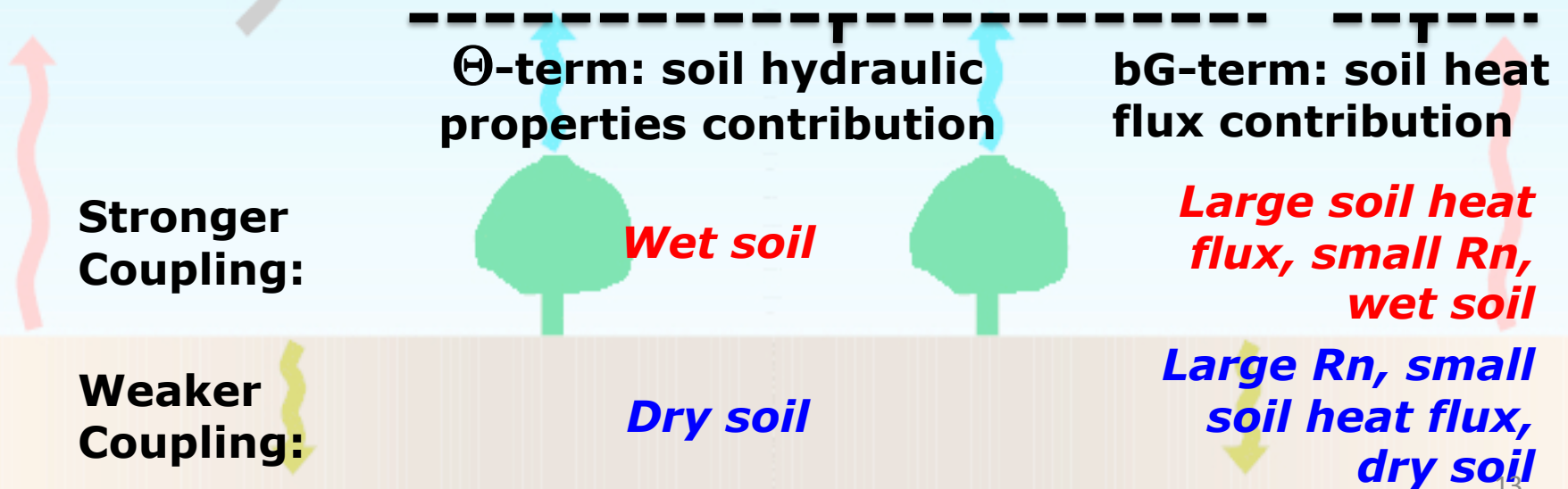
Evap. frac. for bare soil evaporation:
(can't use $E_d = f(\Theta)E_p$)

$$ef_d = \frac{\rho_w L_v}{R_n - G} \left[\frac{\delta\Theta_{ns}}{\delta z} D_\Theta + K_\Theta \right]$$

Mahrt & Pan (1983)

Evap. fraction change with soil moisture change:

$$\frac{\partial \ln ef_d}{\partial \Theta} = \frac{1}{\Theta_{ns}} \left\{ \frac{[\Theta_{ns} + (\beta + 2)\delta\Theta_{ns}] s_\Theta + (2\beta + 3)}{1 + \delta\Theta_{ns} s_\Theta} + \frac{b\beta G}{R_n - G} \right\}$$



Near-surface Interactions:

“Bare” soil

“ ωb ” coupling parameter =

$\Theta + bG$ terms

Additional factors:

• $T \downarrow$ $R_n \downarrow$ $RH \downarrow \rightarrow \omega b \uparrow$

• forest > grass (roughness/partial veg cover)

• sand > clay

• Θ term >> bG term (generally)

**Stronger
Coupling**

**Weaker
coupling**

Fluxnet Data Sets?

strong



TURBULENCE

weak

dry



SOIL MOISTURE

wet

Land-PBL Interaction: Cloud Formation

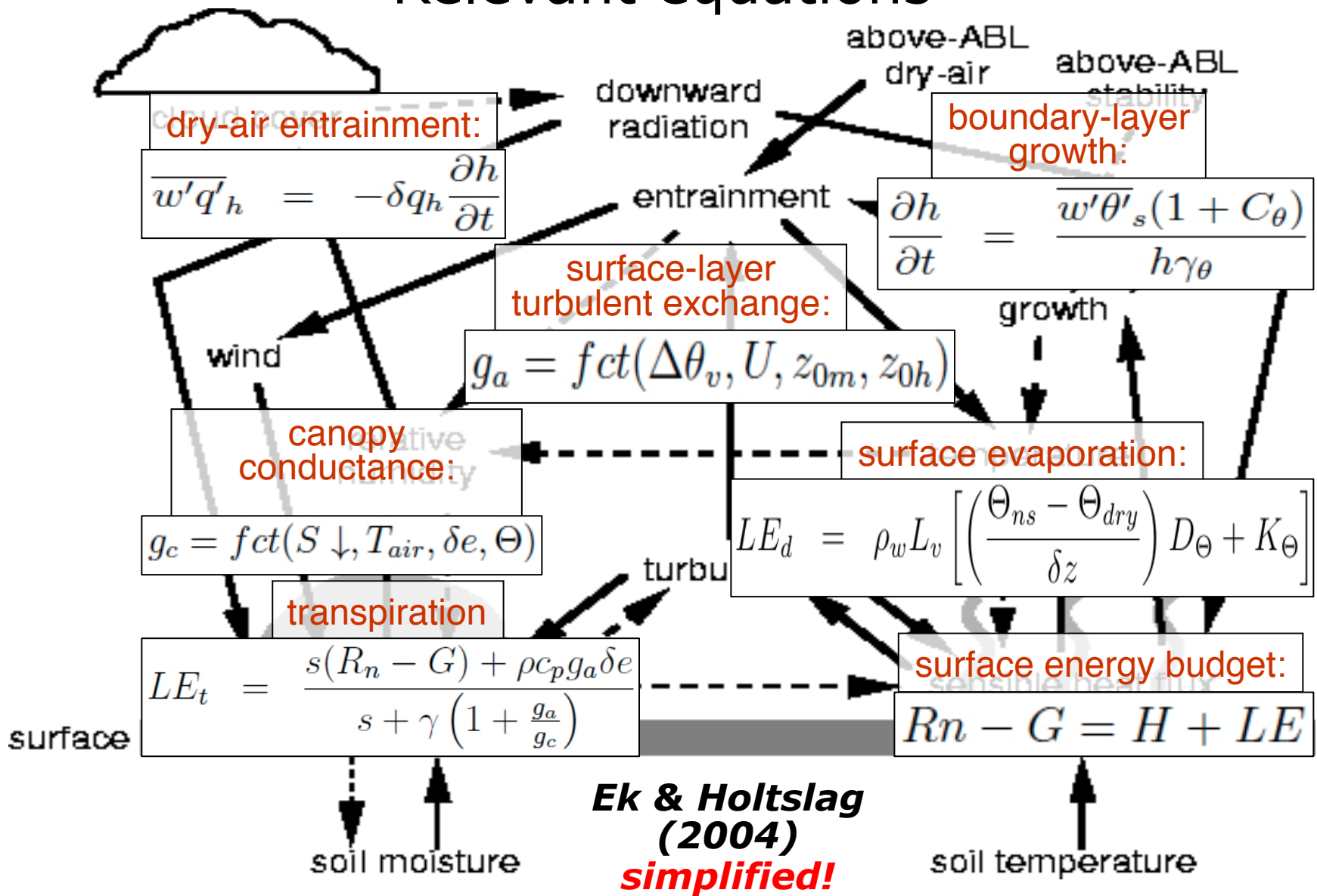
- How does land surface impact onset of Cumulus?
- What is the role of soil moisture and atmospheric processes?
- How to quantify?
- Relative humidity evolution (RH tendency) at the Atmospheric Boundary Layer (ABL) top is expected to control cloud initiation:

$$RH = q/q_s$$

q = specific humidity (g/kg)
 q_s = saturation specific humidity (g/kg)

Land-PBL Interaction:

Relevant equations



Land-PBL Interaction:

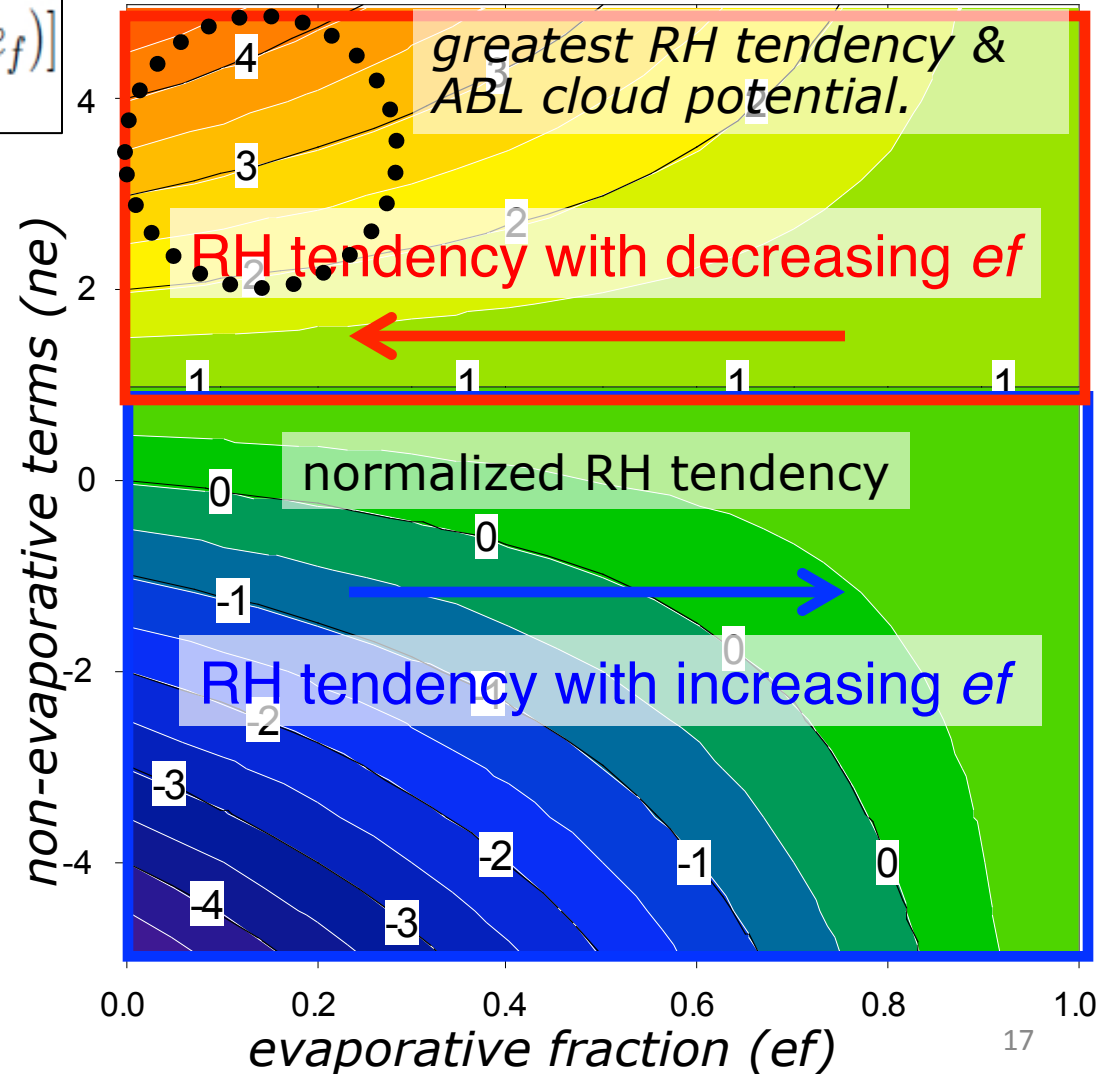
RH tendency at PBL top

$$\frac{\partial RH}{\partial t} = \left(\frac{R_n - G}{\rho L_v h q_s} \right) [e_f + ne(1 - e_f)]$$

$$ne = L_v / c_p (1 + C_\theta) \left[\frac{\Delta q}{h \gamma_\theta} + RH \left(\frac{c_2}{\gamma_\theta} - c_1 \right) \right]$$

- ABL-growth regime (weaker above-PBL stability & Δq small)

- surface moistening regime (stronger above-PBL stability and/or Δq large)



Land-PBL Interaction:

Land & PBL processes and RH tendency

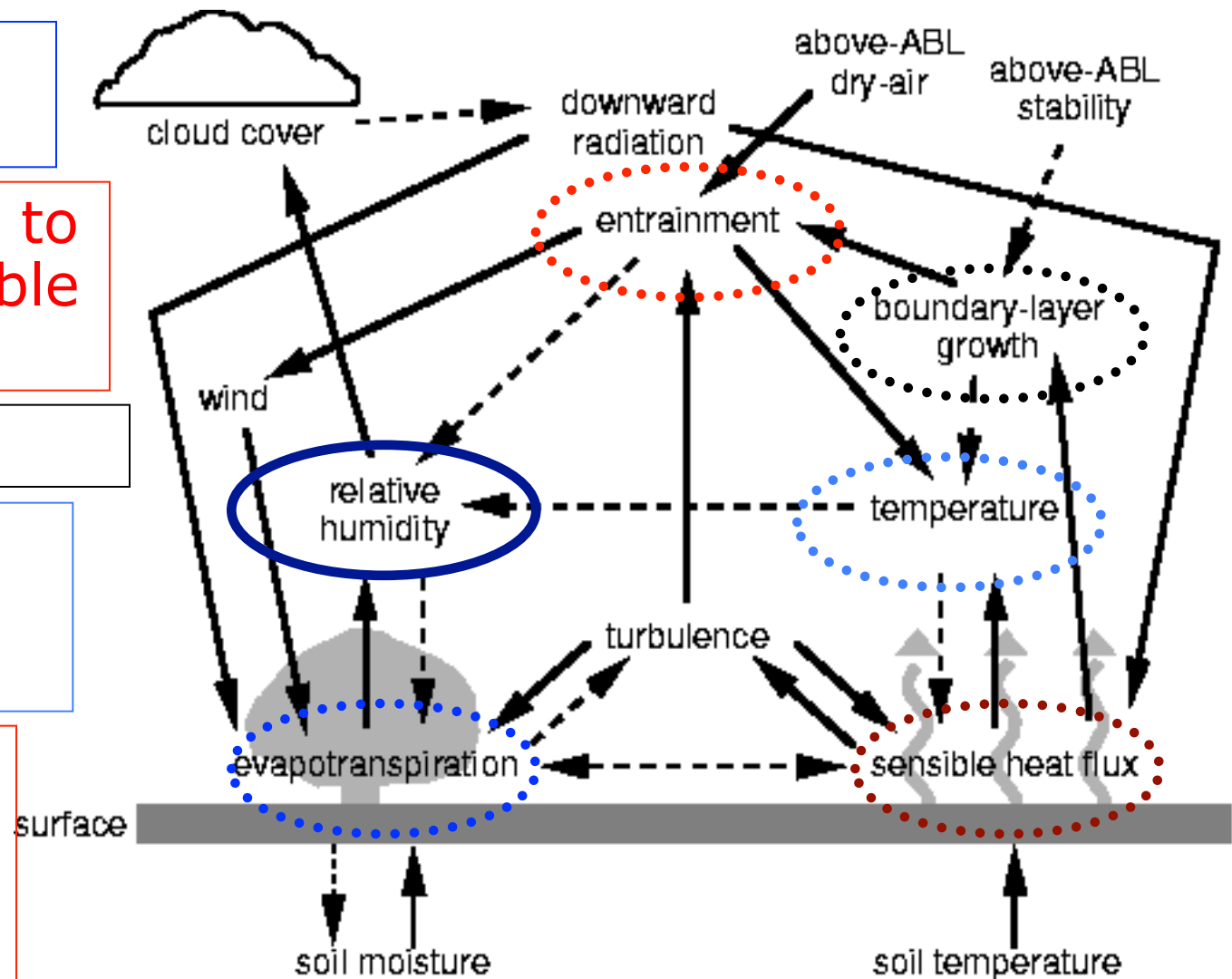
Surface evaporation

Warming due to surface sensible heat flux

ABL growth...

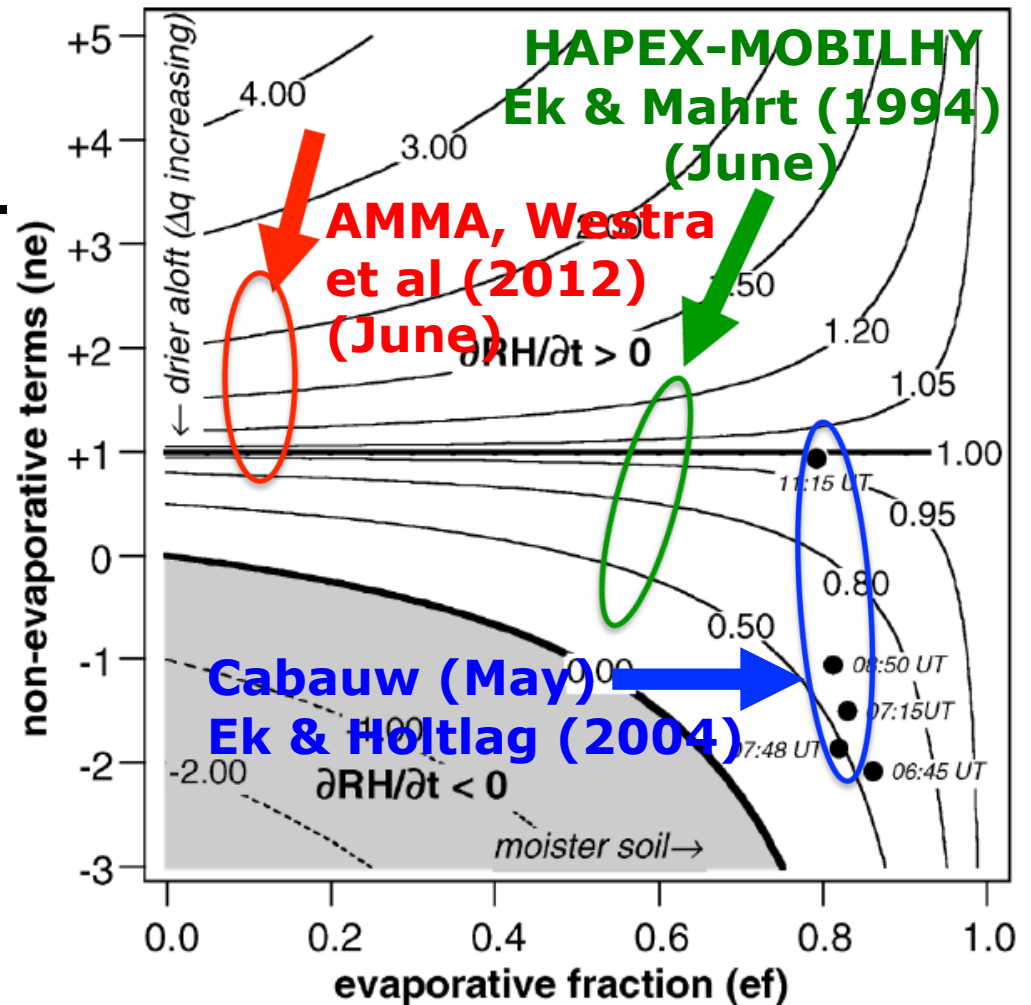
...ABL-top temperature decrease

...ABL-top dry-air & warm-air entrainment



Land-PBL Interaction: Observations

- Evaporative fraction \sim constant, ne increases during day. How general?
- Observed mid-afternoon ABL cloud formation (<20% cover)
- Need to evaluate data sets from more field programs.



Land-PBL Interaction:

Sensitivity tests for coupled land-PBL model and cumulus initiation

**stronger
inversion**

moist air

dry air

**weaker
inversion**

- Examine role of soil moisture: vary soil moisture from dry to wet,
- Vary inversion strength from weak to strong,
- Vary dry air above dry air above the boundary layer from dry to moist,
- Other tests: different vegetation and soil types, background advection, different regions & seasons.

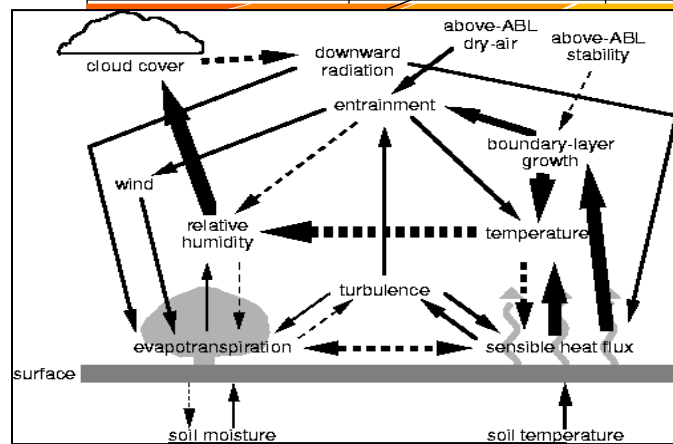
dry soil

wet soil

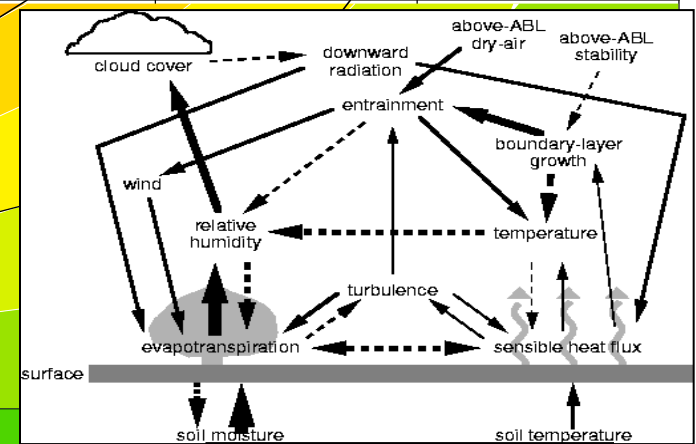
Land-PBL Interaction:

Model sensitivity tests and RH tendency

**WEAK
ABOVE-ABL
STABILITY**

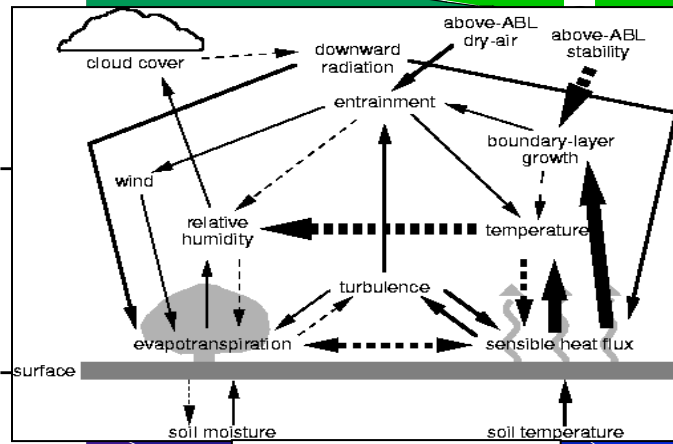


most clouds

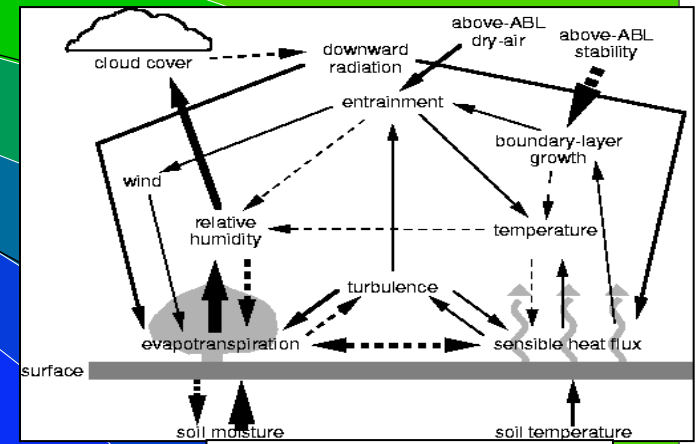


some clouds

**STRONG
ABOVE-ABL
STABILITY**



no clouds



some clouds

0.0

0.2

0.4

0.6

0.8

1.0

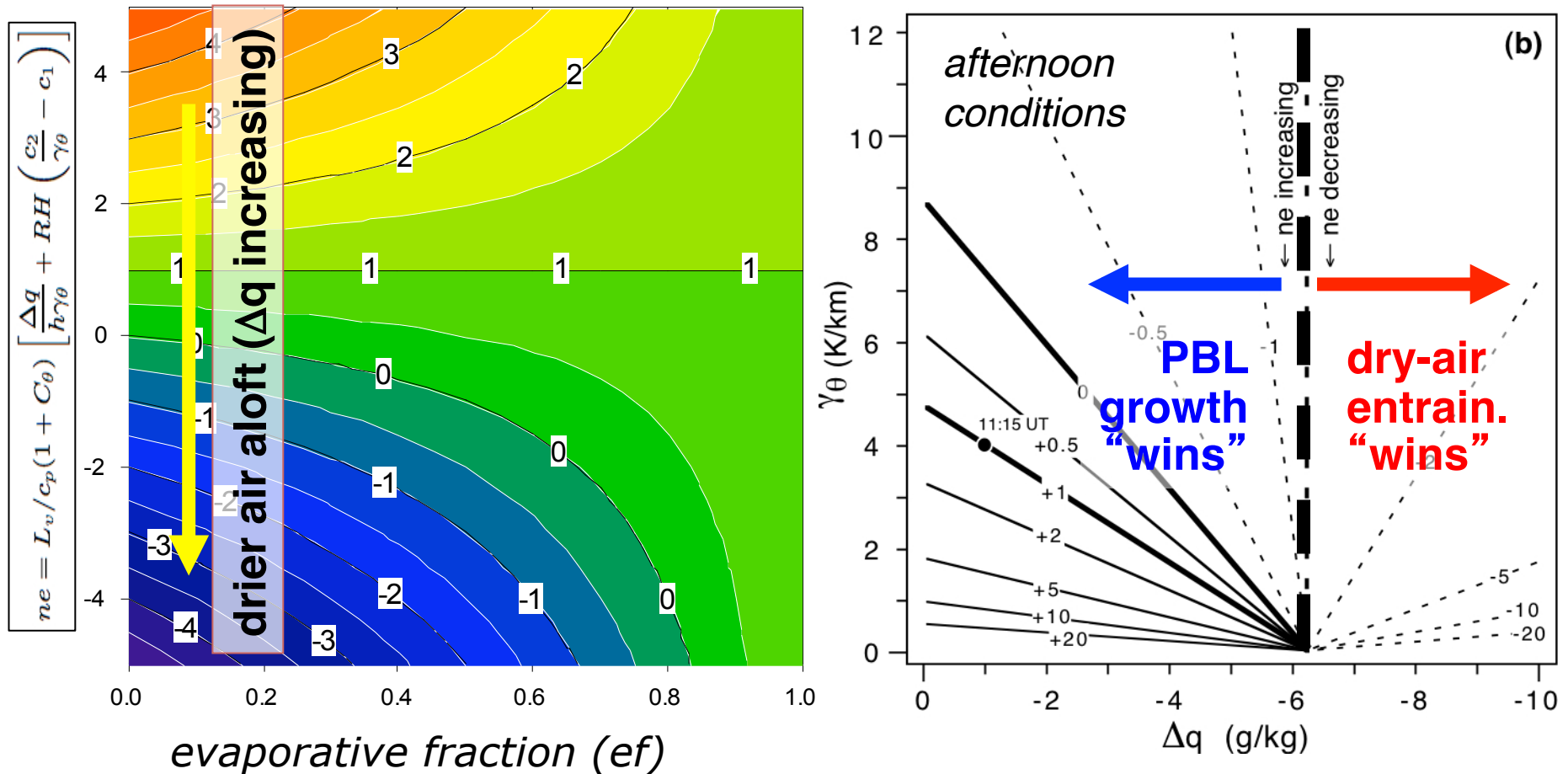
DRY SOIL



MOIST SOIL

Land-PBL Interaction:

Boundary-layer growth vs dry air entrainment



- if $\Delta q >$ critical value (more negative/drier, =fct(h, RH)), then ne decreases with decreasing above-ABL stability, so dry-air entrainment “wins” over BL growth.

Land-PBL Interaction:

Diurnal land-atmos. coupling experiment (DICE)

Objective: Assess impact of land-atmosphere feedbacks.

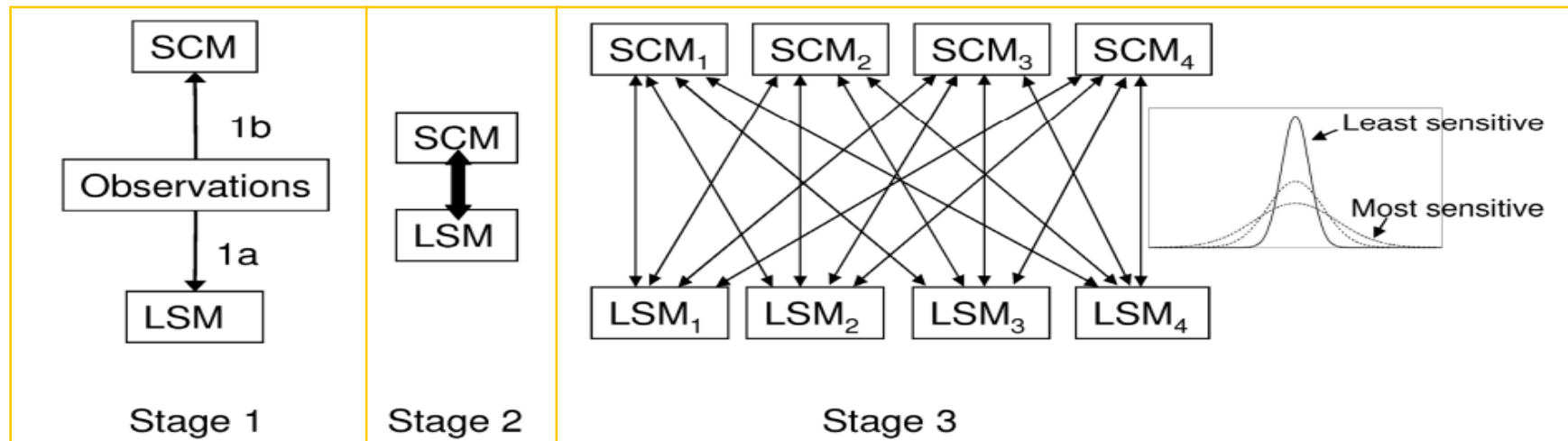
Stage 1: stand alone land, and stand alone single column model (SCM).

Stage 2: Coupled land-SCM.

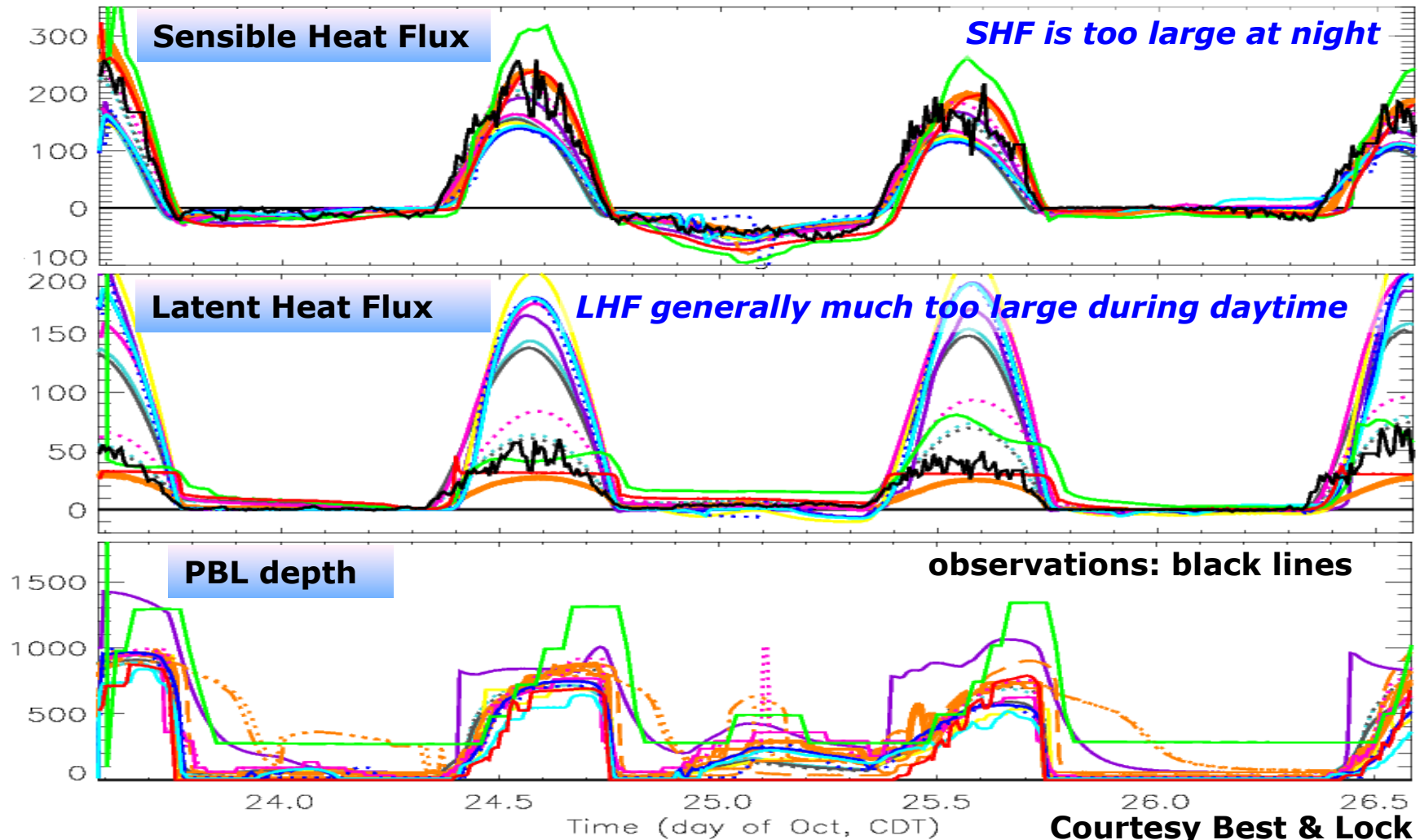
Stage 3: Sensitivity of different LSM and SCM to variations in forcing.

Data Set: CASES-99 field experiment in Kansas, using 3 days: 23-26 Oct 1999, 19UTC-19UTC.

Joint GEWEX GLASS-GASS project –outgrowth of GABLS2 where land-atmosphere coupling was identified as an important mechanism. Lead by Martin Best and Adrian Lock (UKMO).



Land-PBL Interaction: Initial DICE phase 2 (coupled) results

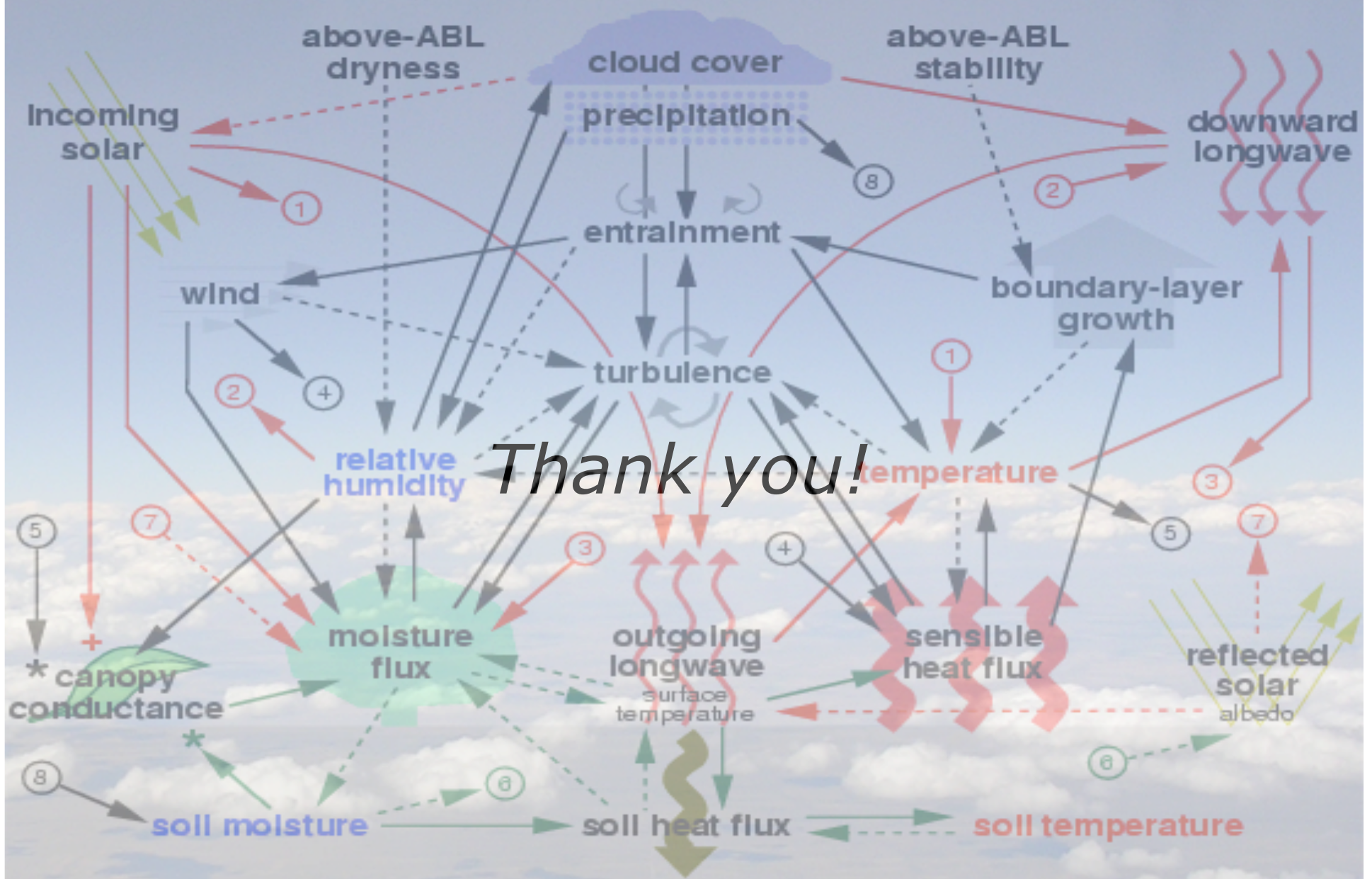


First DICE workshop at UKMO, Exeter, UK, 14-16 Oct 2013.

Summary

- Process level understanding is required to properly represent land-atmosphere interaction (e.g. near-surface, land-PBL) in weather and climate models to get the "*Right answers for the right reasons.*"
- Data wish list: Extensively "mine" **Fluxnet** land data sets and **PBL field programs** for many different regions/seasons (including diurnal cycle); good soil moisture measurements & bare soil sites needed.
- Collaborative efforts from GEWEX GLASS PLUMBER/PALS, LoCo, GLASS-GASS DICE, GABLS, and other programs/projects → use such testing procedures in our model development. "*Step-wise*", "*Pyramid*"
- Important consideration: scale-dependencies and single-site representativeness vs model grid scale.

land-surface - ABL - radiation interactions



Thank you!

+ positive feedback for C3, C4 plants, negative feedback for CAM plants
 * negative feedback above optimal values

→ surface layer/ABL processes
 → land-surface
 → radiation
 - - - → positive feedback
 - - - → negative feedback